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WATER-RESOURCES APPRAISAL OF THE WET MOUNTAIN VALLEY, IN PARTS OF CUSTER AND FREMONT COUNTIES, COLORADO

# U. S. GEOLOGICAL SURVEY



Water-Resources Investigations 78-1

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## UNITED STATES DEPARTMENT OF THE INTERIOR GEOLOGICAL SURVEY

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By Clark J. Londquist and Russell K. Livingston

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Open-File Report

February 1978

## UNITED STATES DEPARTMENT OF THE INTERIOR

CECIL D. ANDRUS, Secretary

## GEOLOGICAL SURVEY

W. A. Radlinski, Acting Director

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## METRIC CONVERSION

Customary units used in this report may be expressed as metric units by use of the following conversion factors:

Multiply customary units	By	To obtain metric units
inch (in.)	25.40	millimeters
foot (ft)	.3048	meter
foot per mile (ft/mi)	.1894	meter per kilometer
mile (mi)	1.609	kilometers
square mile (mi <sup>2</sup> )	2.590	square kilometers
acre	.4047	hectare
acre-foot (acre-ft)	1.233×10 <sup>-3</sup>	cubic hectometers
gallon per minute (gal/min)	.06309	liter per second
cubic foot per second (ft <sup>3</sup> /s)	.02832	cubic meter per second
cubic foot per second per square	.01093	cubic meter per second per
mile $[(ft^3/s)/mi^2]$		square kilometer

WATER-RESOURCES APPRAISAL OF THE WET MOUNTAIN VALLEY, IN PARTS OF CUSTER AND FREMONT COUNTIES, COLORADO

By Clark J. Londquist and Russell K. Livingston

#### ABSTRACT

The Wet Mountain Valley is an intermontane trough filled to a depth of at least 6,700 feet with unconsolidated deposits. Ground water occurs under both artesian and water-table conditions within the basin-fill aquifer and groundwater movement is toward Grape and Texas Creeks. The depth to the water table is less than 10 feet in an area of about 40 square miles along the central part of the valley and is less than 100 feet in most of the remainder of the valley. Ground water stored in the upper 100 feet of saturated basin-fill sediments is estimated to total 1.5 million acre-feet.

Yields greater than 50 gallons per minute generally can be expected from wells in the central part of the basin-fill aquifer, and yields less than 50 gallons per minute are generally reported from wells around the edge of the basin-fill aquifer. Yields of wells in the mountainous areas are generally less than 10 gallons per minute.

The 540-square-mile study area is drained by two principal streams, Grape Creek and Texas Creek. Most streamflow occurs as a result of snowmelt runoff during June and July. In addition to a stream-gaging station on Grape Creek near Westcliffe, which has operated since 1924, there are two continuous-record and four partial-record stations in the study area that were established in 1974. The long-term annual runoff at these seven stations ranges from an estimated 0.02 cubic foot per second per square mile to an estimated 1.17 cubic feet per second per square mile, generally increasing with station altitude.

DeWeese Reservoir, just outside of the study area, has a design storage capacity of 1,772 acre-feet. All other reservoirs in the area have a combined design storage capacity of about 215 acre-feet.

Generalized annual water budgets for two areas in the Wet Mountain Valley indicate that surface-water outflow is only 7 to 11 percent of the total water supply from precipitation and other sources. The remaining water is lost to the atmosphere by evapotranspiration. Ground-water underflow and changes in ground-water storage are considered insignificant in both water budgets.

The quality of both the surface and ground water is generally within the recommended limits for drinking water set by the U.S. Public Health Service. The hardness, as calcium carbonate, of the ground water is generally greater than 100 milligrams per liter.

#### INTRODUCTION

The Southeastern Colorado Water Conservancy District entered into a cooperative agreement with the U.S. Geological Survey in 1970 to conduct a water-resources investigation of the Wet Mountain Valley (fig. 1). The purpose of this investigation was to determine the quantity and quality of both the surface and the ground water within the valley. This information will be used in the management of the water resources of the area.

Westcliffe and Silver Cliff are the only towns with centralized populations in the area, each with a population of less than 400 people. The economy of the region is based primarily on agriculture. Irrigation plays a major role in crop production throughout the valley, particularly in the valley floor north and south of Westcliffe.

The scope of this investigation included the determination of the quality of water in the area, the amount of recoverable water stored in the first 100 ft of saturated alluvial material, expected yields from properly constructed wells, depths to water in the alluvial aquifer, estimates of surface-water flow from small ungaged drainage basins, and a water budget for the Grape Creek drainage area. Streamflow was measured at 7 sites, water samples for chemical analysis were collected from 20 wells and at 2 surfacewater sites, and a network of 31 wells was measured bimonthly during 1970-75 (p1. 1).

The authors wish to acknowledge the cooperation and assistance provided by the many landowners in the area, particularly those who permitted access to their wells and property for the many measurements and tests necessary for the successful completion of this investigation.

#### GEOHYDROLOGIC SETTING

The Wet Mountain Valley is an intermontane trough filled with unconsolidated deposits of Cenozoic age, bordered by mountain ranges composed of volcanic rocks of Tertiary age and igneous, metamorphic, and sedimentary rocks of pre-Tertiary age (MacNish, 1966). The approximate location and extent of the various geologic units within the study area are shown on figure 2.

The study area includes about 540  $\text{mi}^2$  and is primarily drained by two north-flowing tributaries of the Arkansas River. Grape Creek, the principal stream, drains the southern 320  $\text{mi}^2$  of the study area. Texas Creek, the other large stream in the study area, drains an area of 144  $\text{mi}^2$ . The remainder of the area is drained by several small streams that flow directly into the Arkansas River west of Texas Creek.

Precipitation generally increases with land-surface altitude in the Wet Mountain Valley. Based on the areal distribution of the average annual precipitation from 1931 to 1940 (fig. 3) (U.S. Weather Bureau, 1967),





Figure 2 .-- Generalized geologic map.



Figure 3.--Distribution of average annual precipitation 1931-60 (from U.S. Weather Bureau, 1967).

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precipitation ranges from less than 16 in. per year on the valley floor (altitude about 8,000 ft) to more than 40 in. per year (altitude about 13,000 ft) along the Sangre de Cristo Range. About 37 percent of the total annual precipitation in the area occurs in the form of snow.

#### GROUND-WATER OCCURRENCE

Ground water occurs in two different hydrologic environments within the study area--the mountainous area, which is composed chiefly of consolidated rocks, and the valley-floor area, which is underlain by unconsolidated basinfill material. The availability of ground water in the Wet Mountain Valley is summarized in figure 4.

#### Mountainous Area

The Sangre de Cristo Range, on the west side of the valley, consists of igneous, metamorphic, and complexly folded and faulted sedimentary rocks of pre-Tertiary age. The Wet Mountains, on the east side of the valley, are composed of igneous and metamorphic rocks of pre-Tertiary age which are overlain by Tertiary volcanic rocks at some localities.

In the mountainous area, ground water is chiefly contained in the fractured and weathered zones of the consolidated rocks or in minor, isolated pockets of overlying glacial and alluvial material. The weathered zones of the consolidated rocks are generally limited to the upper few tens-of-feet of the rock and the fractures generally become tighter and, hence, contain less water with depth. Yields from wells completed in consolidated rock are unpredictable due to the uncertain distribution of fractures and thickness of the weathered zones but, generally, are less than 10 gal/min (fig. 4). Yields greater than 10 gal/min might be obtained where the fractures or weathered zones are in direct hydraulic connection with the water in nearby streams.

The glacial and alluvial material in the mountainous area is generally restricted to the canyons and is not widespread. Where these materials occur above the water level of nearby streams, they are generally well drained and can be expected to yield less than 10 gal/min. In areas where the unconsolidated materials are in direct hydraulic connection with the water in nearby streams, well yields greater than 10 gal/min may be obtained.

In the volcanic rocks along the eastern margin of the valley, reported well yields vary considerably. In areas where the volcanic rocks are tuffaceous and of low permeability, the reported well yields are relatively small. Several wells have reportedly been drilled to depths ranging from 200 to 300 ft and have yields of less than 2 gal/min . Locally, where rhyolitic rock is encountered, well yields of as much as 30 gal/min have been reported from wells of less than 50 ft deep.



Figure 4.-- Availability of ground water.

#### Valley-Floor Area

The basin-fill deposits in the valley-floor area cover an area of approximately 233 mi<sup>2</sup> and are the principal source of ground water in the study area. The basin-fill material is composed primarily of alluvial material derived during Tertiary time from the mountains surrounding the valley. Glaciation of the Sangre de Cristo Range has formed glacial moraine and outwash deposits which are superimposed on the Tertiary basin-fill material in some localities. Since the disappearance of the glaciers from the area, the trough has continued to slowly fill with alluvial material derived from the surrounding mountains.

An electrical resistivity survey was made from west to east across the valley about 3 mi south of Westcliffe (pl. 1). The resistivity section (fig. 5) indicates that the thickness of the basin-fill material ranges from about 1,000 ft at the east end of the section to about 6,700 ft at the west (Zohdy and others, 1971). A test hole, SC 23-72- 6AAB1 (the system of end numbering wells is explained in the Supplemental Information section of this report), was drilled along the line of the section, about 0.5 mi west of Grape Creek, to a depth of 1,200 ft. The geologic log of this test hole is included as Supplemental Information in this report. The log indicates that the upper 1,200 ft of the basin-fill material, in this area, is composed primarily of sand and gravel with interbedded and intergranular clay. The resistivity section indicates that the basin-fill material contains increasing amounts of clay with depth, or that the water in the basin-fill material contains increasing amounts of dissolved solids with depth.

Ground water occurs under both artesian and water-table conditions within the basin-fill aquifer. Artesian conditions are created by interbedded and intergranular clays confining the water beneath them. These confining layers probably are not continuous or completely impermeable and there is most likely a considerable amount of leakage between the confined and unconfined parts of the aquifer. The artesian conditions exist, to some extent, throughout most of the area of the basin fill. Along the central part of the valley, there are several flowing wells (pl. 1).

The depth to the water table is less than 10 ft in an area of about 46 mi<sup>2</sup> along the central part of the valley, as shown on figure 6. The depth to the water table is less than 100 ft in most of the remainder of the valley except for the extreme southern and southeastern parts of the valley, where it is generally greater than 100 ft.

There is an estimated 1.5 million acre-ft of recoverable ground water stored in the upper 100 ft of saturated basin-fill material. Hydrographs from selected wells are shown in figure 7. The hydrographs indicate that for the observed period of record, 1970-75, the water levels are undergoing only seasonal fluctuations and that no changes in the amount of ground water in storage are taking place.





Figure 6.-- Depth to water in the basin-fill aquifer.





The basin-fill area has been divided into two areas on the basis of reported well yields (fig. 4). In the central part of the valley, wells of adequate depth, diameter, and screen length can be expected to have sustained yields of at least 50 gal/min. Several wells in this area have reported yields greater than 300 gal/min. (See Records of Selected Wells in the Supplemental Information section of this report.)

#### GROUND-WATER MOVEMENT

Water enters the aquifers in the mountainous area as infiltration from precipitation or streams. The water then moves downgradient where it is discharged by springs to the numerous small streams, lost to evapotranspiration, or discharged into the basin-fill aquifer as underflow.

Extensive faults have been mapped along both the east and west sides of the basin between the basin-fill material and the consolidated rock (Scott and Taylor, 1975). The low permeability fault gouge associated with these faults probably serves as at least a partial barrier to water movement between the consolidated rock and the basin-fill material. This barrier and the general low permeability of the consolidated rock limit the direct discharge of water from the consolidated rock to the basin-fill material to very small amounts. Water enters the basin-fill aquifer as infiltration from precipitation falling directly on the aquifer or as rim inflow from the surrounding mountainous area.

The basin-fill material is surrounded, except for the southern end, by bedrock of low permeability. The volcanic rocks in the basin (fig. 2) are not continuous and do not serve as a conduit for transporting water out of the basin. Once the water enters the basin-fill aquifer, it moves toward the center of the basin and does not flow out of the basin to the south under Promontory Divide. Very little, if any, water leaves the basin-fill aquifer as underflow. The direction of ground-water flow in the basin-fill aquifer is shown on plate 1. Ground water moves from higher to lower potentiometric levels. As illustrated by the potentiometric map (pl. 1), ground water moves from the boundaries of the basin-fill aquifer toward Grape and Texas Creeks.

Water leaves the aquifer in several ways. Direct evapotranspiration from the water table occurs in the area where the depth to water is less than 10 ft (fig. 6). Water is discharged from the aquifer into Grape Creek along the reach from just south of Westcliffe to the edge of the basin fill and to Texas Creek in the reach just before it leaves the basin-fill material. Some water also is lost to small streams that cross the northern end of the valley-fill material and then discharge directly into the Arkansas River. Water isdischarged through wells either by pumping or by artesian flow. Part of the water discharged from wells infiltrates the land surface and percolates back down to the water-table aquifer, while the remainder is consumed, lost to evapotranspiration, or runs off into the streams.

#### SURFACE WATER

The surface-water resources of the Wet Mountain Valley are vital to the agricultural economy of the area. Most of the water for irrigation is obtained by diversion of surface water rather than pumping of ground water. A tabulation of water rights on Grape Creek and Texas Creek is presented in table 1.

Observed streamflow consists of runoff from precipitation, return flow from irrigation, seepage from ground water, and (or) water released from reservoir storage. A hydrograph of mean monthly discharges (1925-75) of Grape Creek near Westcliffe (station 07095000; the station numbering system is explained in the Supplemental Information section of this report) is shown in figure 8. Runoff from melting snow dominates the hydrograph during May, June, and July. Streamflow during the low-flow or base-flow period in fall and winter is due largely to irrigation return flows and ground-water discharge.

The streamflow hydrograph shown in figure 8 does not indicate the effects of significant surface-water storage in the watershed. A tabulation of lakes and reservoirs in the study area and their location, design surface area, and design storage capacity is given in table 2. DeWeese Reservoir, the largest reservoir in the vicinity of the study area, is located about 0.5 mi downstream from station 07095000. The water-storage capacity of this reservoir is 1,772 acre-ft. The remaining five reservoirs in the study area have a combined water-storage capacity of about 215 acre-ft.

#### Streamflow [Variable]

The Wet Mountain Valley has large year-to-year variations in streamflow. A hydrograph of annual runoff recorded by the Grape Creek near Westcliffe gage (07095000) is shown in figure 9. From 1924 to 1975, annual recorded runoff ranged between 5,120 acre-ft in 1963 and 78,850 acre-ft in 1942. As previously stated, the study area has only about 215 acre-ft of available water storage in surface-water reservoirs.

Other than at the Grape Creek near Westcliffe (07095000) gaging station that was established in 1924, very few records of streamflow have been collected in the study area. The streamflow-gaging station on Grape Creek near Canon City (07095500) was operated during 1907-09 and the station on Texas Creek at Texas Creek (07094000) was operated during 1923. Unpublished records of streamflow exist for Dry Creek near Westcliffe during the late 1960's and for Texas Creek at Texas Creek since September 1970.

In order to improve areal coverage of streamflow information, two continuous-record and four partial-record gaging stations were installed in August 1974 as part of this study. These stations are listed in table 3 and located on plate 1. Data obtained from these six additional stations were useful in evaluating the streamflow characteristics of the many smaller streams draining the perimeter of the study area.

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	Location <sup>2</sup>			Amount of decree,	
Name of diversion	Town- ship	Range	Section	in cubic feet per second	Date of decree
	G	RAPE CI	REEK		
Aldrich No. 2	225	072W	32	0.34	04-15-1870
Voris Bros. No. 1	21S	073W	25	2.34	07-31-1871
DeWeese Dye Main	19S	070W	06	2.34	07-31-1871
Voris Bros. No. 1	21S	073W	25	2.34	07-31-1871
Pinto	235	072W	21	.50	05-01-1872
Southern	235	072W	27	1.38	06-12-1873
Voris Bros. No. 1	21S	072W	25	•95	10-31-1873
DeWeese Dye Main	19S	070W	06	•95	10-31-1873
Voris Bros. No. 1	215	072W	25	.95	10-31-1873
Difz	235	072W	27	.83	06-15-1874
Rifster No. 2	23S	072W	34	.88	10-10-1874
С В Н	23S	072W	34	.50	10-10-1874
Pinto	23S	072W	21	2.50	03-03-1875
Gordon W. Smith No. 1	22S	072W	19	3.00	05-15-1875
Elze No. 3	24S	072W	04	.69	05-31-1875
Elze No. 5	24S	072W	03	.16	05-31-1875
Voris Jarvis No. 1	22S	073W	01	2.61	05-31-1875
Risser and Locke	235	072W	05	5.11	06-01-1875
Hulmuth	23S	072W	16	2.50	10-05-1875
R B	235	072W	20	.89	05-05-1876
Chetelate No. 2	23S	072W	20	•33	06-01-1876
Schulz No. 3	24S	072W	03	.66	06-01-1876
С В Н	46N	012E	13	1.30	06-05-1876
Colfax	23S	072W	04	2.33	06-10-1876
Voris Bros. No. 2	225	073W	01	1.33	06-30-1877
DeWeese Dye Main	195	070W	06	1.33	06-30-1877
Voris Bros. No. 2	22S	073W	01	1.33	06-30-1877
Schoolfield No. 4	22S	072W	05	.56	04-01-1878
Schulz No. 5	24S	072W	03	.33	05-15-1878
Voris Bros. No. 1, 2d Enl	215	073W	25	1.09	10-31-1879
DeWeese Dye Main	195	070W	06	1.09	10-31-1879
Voris Bros. No. 1, 2d Enl	215	073W	25	1.09	10-31-1879
Total				- 44.53	

Table	1Surface-w	ater d	iver	sion :	rights	prior	to	1880
	along	Grape d	and	Texas	Creeks	,1		

		Locatio	on <sup>2</sup>	Amount of decree,	
Name of diversion	Town- ship	Range	Section	in cubic feet per second	Date of decree
	Т	EXAS C	REEK		
Hill No. 1	47N	012E	12	1.05	08-31-1869
Hill No. 2	47N	012E	12	1.05	08-31-1869
Pitter No. 1	21S	073W	29	1.00	04-30-1871
Likely & McCormick	47N	012E	14	1.32	06-01-1871
McCormick	47N	012E	13	1.00	06-01-1871
Mill				1.00	09-01-1871
Ritter No. 2	21S	073W	29	1.00	05-01-1872
Duckett No. 1	47N	012E	26	1.00	05-10-1872
Belknap & Hendrickson	47N	012E	35	1.00	05-10-1872
Duckett No. 1	47N	012E	26	1.00	05-10-1872
Belknap & Howard	46N	012E		2.76	05-31-1872
Dissmore No. 2 McClurkens	21S	073W	31	1.00	07-24-1872
McClurken No. 1	21S	073W	31	1.00	07-30-1872
Hugg No. 3	21S	073W	32	1.06	11-30-1872
Hugg No. 4	21S	073W	32	1.00	11-30-1872
Duckett No. 2	47N	012E	26	1.32	05-15-1873
Hill Ditch No. 2	47N	012E	12	1.05	01-17-1874
Myers	215	073W	16	1.00	05-01-1874
J. M. Duckett	47N	012E	34	3.50	05-31-1874
Burgman No. 2	47N	012E	33	1.00	04-30-1875
Hill No. 1	47N	012E	12	1.05	04-14-1876
Likely	47N	012E	12	1.60	03-06-1877
Belknap & Hendrickson	47N	012E	35	1.58	05-31-1878
Total				- 29.34	

# Table 1.--Surface-water diversion rights prior to 1880 along Grape and Texas Creeks<sup>1</sup>--Continued

 $^1 \mbox{Source:}$  Colorado Division of Water Resources, Office of the State Engineer.  $^2$ Refers to location of point of diversion.





		Locatio	on <sup>2</sup>		Design	Design	
Name	Town- ship	Range	Section	Owner	area, in acres	in acre-feet	
Balman Reservoir	46N	12E	6	U.S. Forest Service	3.6	50.9	
Hook Reservoir	46N	12E	2	Thomas Hook	2.0	35.0	
DeWeese Reservoir-	215	72W	20	DeWeese Dye Ditch and Reservoir Co.	148	1,772	
Lake of the Clouds Reservoir No. 1-	45N	12E	(21)	U.S. Forest Service	15.0	48.0	
Lake of the Clouds Reservoir No. 2-	45N	12E	(21)	do	15.0	36.0	
Lake of the Clouds Reservoir No. 3-	45N	12E	(21)	do	(15)	(45)	
<sup>1</sup> Source: Colo	orado	Divis	ion of N	Water Resources, O	ffice of	the State	

Table 2.--Principal lakes and reservoirs in the Wet Mountain Valley and vicinity  $^{\rm l}$ 

Engineer. <sup>2</sup>Refers to location of dam or outlet.

Table 3.--Stream-gaging stations in

Stream	U.S. Geolog-			Location			
site number <sup>1</sup>	station number	Station name	Operating agency	Latitude	Longitude		
1	07093980	Lake Creek near Hillside, Colo-	U.S. Geological Survey	38°15'49''	105°39'41''		
2	07094000	Texas Creek at Texas Creek (near Cota- paxi), Colo	Office of the Colorado State Engineer	38°24'30''	105°35'00''		
3	07094530	Grape Creek near Bradford, Colo-	U.S. Geological Survey	37°57'03''	105°27'03''		
4	07094600	South Colony Creek near Westcliffe, Colo	do	37°59'57''	105°29'25''		
5	07094700	Antelope Creek near Rosita, Colo	do	38°01'18''	105°21'43''		
6	07094800	Venable Creek near West- cliffe, Colo	do	38°05'02''	105°33'52''		
7	07094900	Middle Taylor Creek near Westcliffe, Colo	do	38°06'30''	105°36'03''		
8	07095000	Grape Creek near Westcliffe, Colo	Office of the Colorado State Engineer	38°11'10''	105°28'59''		

 $^1{\rm Refers}$  to sites shown on figure 2.  $^2{\rm Refers}$  to available record from which flow characteristics were

Drainage area (mi <sup>2</sup> )	Period of record <sup>2</sup>	Mean annual discharge (ft <sup>3</sup> /s)	Maximum observed discharge (ft <sup>3</sup> /s)	Minimum observed discharge (ft <sup>3</sup> /s)	Remarks
9.07	June 1974- September 1976		20	1.6	Partial-record sta- tion, discharge measurements only.
144	April-November 1923		2,800		Records only for 1923 are published.
6.72	June 1974- September 1976		19	.3	Partial-record sta- tion, discharge measurements only.
6.03	October 1974- September 1976	6.65	47	.20	
35.3	June 1974- September 1976		9.4	0	Do.
2.55	June 1974- September 1976		17	. 04	Do.
3.19	October 1974- September 1976	4.10	38	. 40	
320	October 1924- September 1961; October 1962- September 1976	32.3	7,460	. 01	

the Wet Mountain Valley and vicinity

determined; records may exist for other periods.





#### Runoff Analysis

The results of an analysis of runoff in the study area are summarized in table 4. The 1975-76 annual runoff for the four partial-record stations was determined by correlation techniques using the two continuous-record stations and is expected to be within 10 percent of the actual runoff (Riggs, 1969). Observed annual runoff in the study area was generally greatest during the 1975 water year.

Long-term annual runoff (1931-60), also given in table 4, was estimated for the six short-term gaging stations based on the 1975-76 average annual runoff, channel-geometry techniques (Hedman and others, 1972), and regression analysis of basin characteristics (Livingston, 1970). Equations for predicting the long-term mean annual runoff using channel-geometry measurements have a standard error of estimate of 19 percent; those for basin characteristics, (The standard error of estimate is the percentage above or below 47 percent. calculated values within which about two-thirds of the actual values are expected to fall.) The recorded 1975-76 average annual runoff for Grape Creek near Westcliffe is in close agreement with the long-term (1931-60) runoff. Consequently, the 1975-76 average annual runoff for the six short-term gaging stations, along with the determinations made using channel-geometry and regression methods, were averaged to estimate the long-term (1931-60) mean annual runoff at these stations. These data were used in the water-budget analysis which is discussed in a later section of this report.

Annual runoff expressed as a function of drainage area, in cubic feet per second per square mile, generally increases with altitude in the Wet Mountain Valley, as shown in table 4. Values of estimated mean annual runoff for 1931 through 1960 range from about  $0.02 \ (ft^3/s)/mi^2$  for Antelope Creek (altitude 8,290 ft) to about  $1.17 \ ft^3/s)/mi^2$  for Middle Taylor Creek(altitude 9,960 ft). Such data can be used to give an estimate of annual runoff at ungaged sites in the study area. For example, the data shown in table 4 indicate that the annual runoff for a stream with a 5-mi<sup>2</sup> drainage area at an altitude of 9,000 ft is about 5  $ft^3/s$  or  $1.0 \ (ft^3/s)/mi^2$ .

#### Magnitude and Frequency of Floods

Occasional flooding in the Wet Mountain Valley results from either intensive rainstoms or unusual snowmelt-runoff conditions. Information on the magnitude and frequency of floods can be used by planners and designers to minimize damages and hazards in areas subject to flooding.

Where adequate records of flood discharge are available, a relation can be developed between flood magnitude and frequency of occurrence. The log-Pearson Type III distribution has been established as the method for uniform flood-frequency analyses by all Federal agencies (U.S. Water Resources Council, 1976). In order to reduce the effects of possible time-sampling errors associated with such analyses, a weighting procedure can be used (Sauer, 1974; McCain and Jarrett, 1976). For example, the 100-year flood discharge for Grape Creek at Westcliffe (07095000) is 4,900 ft<sup>3</sup>/s by log-

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Table 4.--Annual runoff in the

Station nome	Drainage	Alti-	Observed an	nual runoff, in acre-feet		
(station number)	square miles	in feet	1975 Water year	1976 Water year	1975-76 Average	
Lake Creek near Hillside (07093980) <sup>4</sup> -	- 9.07	8,300	3,590	3,610	3,600	
Grape Creek near Bradford (07094530) <sup>4</sup> -	- 6.72	8,690	3,330	2,960	3,140	
South Colony Creek near Westcliffe (07094600)	- 6.03	8,940	5,060	4,570	4,820	
Antelope Creek near Rosita (07094700) <sup>4</sup>	- 35.3	8,290	780	50	420	
Venable Creek near Westcliffe (07094800)	<sup>4</sup> 2.55	9,020	2,160	1,510	1,840	
Middle Taylor Creek nea Westcliffe (07094900)	r - 3.19	9,960	3,700	2,240	2,970	
Grape Creek near Westcliffe (07095000)	- 320	7,790	33,710	14,780	24,240	

<sup>1</sup>Hedman, Moore, and Livingston, 1970.

<sup>2</sup>Livingston, 1970.

<sup>3</sup>Figures in parentheses were determined as the mean of the observed regression methods (see text).

<sup>4</sup>Partial-record station.

<sup>5</sup>Value not used in determining mean annual runoff.

Pearson Type III analysis and 9,370  $ft^3/s$  by the weighting procedure. A complete summary of the flood characteristics for Grape Creek at Westcliffe (07095000), including the equations and parameter values used in their determination, is given in table 5.

For natural-flow streams where adequate records of flood discharge are not available, the flood characteristics can be estimated using regional relations given in a manual developed by McCain and Jarrett (1976). The method presented in the manual involves the use of basin characteristics and regionalized regression equations. A summary showing the use of this method to determine the flood characteristics at two partial-record stations, Antelope Creek near Rosita (07094700) and Venable Creek near Westcliffe (07094500) is given in table 5.

Estimated annual runoff, in acre-feet		Mean annual runoff (1931-60 calendar years) <sup>3</sup>			
Channel-geometry method <sup>1</sup>	Regression method <sup>2</sup>	Acre-feet	Cubic feet per second per square mile		
3,560	<sup>5</sup> 5,170	(3,600)	(0.55)		
<sup>5</sup> 1,120	2,880	(3,000)	(.62)		
2,920	5,930	(4,500)	(1.03)		
380		(400)	(.02)		
1,830	1,930	(1,900)	(1.03)		
2,260	2,850	(2,700)	(1.17)		
		22,460	.10		

Wet Mountain Valley

1975-76 average runoff and the estimated runoff by channel-geometry and

## GENERALIZED WATER BUDGET

The movement of water in a particular area may be expressed as a water budget in which an accounting is made of all water entering or leaving the area. The water budget can be expressed in equation form as:

Water in = Water out ± change in storage,

where "water in" includes precipitation and surface- and ground-water inflow, "water out" includes evapotranspiration and surface- and ground-water outflow, and "change in storage" includes both surface- and ground-water storage and is positive for a net increase in storage.

Station name and number	Basin type	Required parameters
Antelope Creek near Rosita (07094700)	Ungaged site mixed popula- tion (plains/ mountain) re- gion	$\begin{array}{l} Q_{T} = \text{discharge for recurrence interval, T,} \\ D_{T} = \text{depth for recurrence interval, T,} \\ A = 35.3  \text{mi}^{2}  (\text{drainage area}) \\  (9.45  \text{mi}^{2},  \text{plains region;} \\ 25.8  \text{mi}^{2},  \text{mountain region}), \\ P = 18  \text{in.} \\ = \text{mean annual precipitation} \\  (\text{mountain region only}), \\ S_{B} = 114  \text{ft/mi}  (\text{basin slope}) \\ S_{S} = 64  \text{ft/mi}  (\text{streambed slope}) \\ R = 0.70  (\text{reduction factor for site} \\  \text{elevation}). \\ \end{array}$
Venable Creek near Westcliffe (07094900)	Ungaged site mountain region	A=2.55 mi <sup>2</sup> (drainage area), P=28 in. (mean annual precipitation).
Grape Creek near Westcliffe (07095000)	Gaged sitemixed population (plains/moun- tain) region	$\begin{array}{l} {}^{Q}_{T}\left(w\right)^{=weighted \ discharge \ for \ recurrence \ interval, T,} \\ {}^{Q}_{T}(s)^{=station \ value \ of \ the \ flood \ for \ recurrence \ interval, T^{1},} \\ {}^{Q}_{T}(R)^{=regression \ value \ of \ the \ flood \ for \ recurrence \ interval, T,} \\ {}^{N=46 \ (number \ of \ years \ of \ station \ data \ used \ to \ compute \ Q_{T}(s)^{}),} \\ E=10 \ (equivalent \ years \ of \ record \ for \ Q_{T}(R)^{}), \\ E=10 \ (equivalent \ years \ of \ record \ for \ Q_{T}(R)^{}), \\ A= \ 320 \ mi^{2} \ (drainage \ area) \ (220 \ mi^{2}, \ plains \ region; \ 100 \ mi^{2}, \ mountain \ region), \\ P=19 \ in., \\ {}^{N=47 \ ft/mi}_{N}_{S}_{S}=29 \ ft/mi. \end{array}$

<sup>1</sup>Method of determination as described by McCain and Jarrett, 1976. <sup>2</sup>Standard error of estimate is the percentage above or below calculated <sup>3</sup>Highest flood characteristic obtained by the two methods used; in this

Regression equation	Standard error of estimate <sup>2</sup> (percent)	Resultant flood characteristic
Plains region:		
$Q_{10} = 144 \ A^{0.528} \ S_{B}^{0.336}$	31	${}^{3}Q_{10} = 1,620 \text{ ft}^{3}/\text{s}.$
$Q_{50} = 891 \ A^{0.482} \ S_{B}^{0.154}$	24	${}^{3}Q_{50} = 3,820 \text{ ft}^{3}/\text{s}.$
$Q_{100}=1,770 A^{0.463} S_{B}^{0.086}$	28	<sup>3</sup> Q <sub>100</sub> = 5,270 ft <sup>3</sup> /s.
Q <sub>500</sub> =5,770 A <sup>0.432</sup>	45	<sup>3</sup> Q <sub>500</sub> =10,660 ft <sup>3</sup> /s.
$D_{100} = 59.3 \text{ s}_{5}^{-0.517}$	21	<sup>3</sup> D <sub>100</sub> =6.9 ft.
Mountain region:		ч. Ч
Q <sub>10</sub> =0.12 A <sup>0.815</sup> P <sup>1.592</sup>	39	
Q <sub>50</sub> =0.91 A <sup>0.795</sup> P <sup>1.110</sup>	37	
Q <sub>100</sub> =1.88 A <sup>0.787</sup> P <sup>0.932</sup>	38	
Q <sub>500</sub> =8.70 A <sup>0.766</sup> P <sup>0.560</sup>	45	
D <sub>100</sub> =1.44 A <sup>0.187</sup> P <sup>0.059</sup>	28	

gaging stations in the Wet Mountain Valley $^{
m 1}$ 

(Equations given above)

$$Q_{T(W)} = \frac{(Q_{T(S)} \times N) + (Q_{T(R)} \times E)}{N + E}$$

(Equations for  $Q_{T(R)}$  and  $D_{100}$  given above)

$$Q_{10} = 52 \text{ ft}^3/\text{s.}$$
  
 $Q_{50} = 77 \text{ ft}^3/\text{s.}$   
 $Q_{100} = 88 \text{ ft}^3/\text{s.}$   
 $Q_{500} = 115 \text{ ft}^3/\text{s.}$   
 $D_{100} = 2.1 \text{ ft.}$ 

 ${}^{3}Q_{10}(W) = 2,780 \text{ ft}^{3}/\text{s.}$   ${}^{3}Q_{50}(W) = 6,740 \text{ ft}^{3}/\text{s.}$   ${}^{3}Q_{100}(W) = 9,370 \text{ ft}^{3}/\text{s.}$   ${}^{3}Q_{500}(W) = 18,800 \text{ ft}^{3}/\text{s.}$  ${}^{3}D_{100}(W) = 10.4 \text{ ft.}$ 

values within which about two-thirds of the actual values are expected to fall. instance, the characteristics from the plains-region determination are used.

Generalized annual water budgets are given in table 6 for the Grape Creek drainage basin above the Grape Creek near Westcliffe gaging station (07095000), and for that part of this drainage basin consisting of basin-fill material (fig. 10). The area of basin-fill drained by Grape Creek represents 167 mi<sup>2</sup> of the 320-mi<sup>2</sup> drainage basin. The budgets are based on calendar years 1931 through 1960.

#### Precipitation

The average annual precipitation falling on the water-budget areas was determined by use of the distribution of mean annual precipitation for 1931-60 shown on figure 3. The isohyetal method of analysis resulted in a 331,000-acre-ft average for the entire drainage area and a 151,000-acre-ft average for the basin-fill area.

#### Surface-Water Inflow and Outflow

There is no surface-water inflow across the Grape Creek drainage divide, such as transmountain diversions. Surface-water inflow to the basin-fill area consists of runoff from the mountainous regions along the perimeter of the For purposes of the water budget, the 153 mi<sup>2</sup> that contribute inflow area. were separated into the two tributary areas shown on figure 10. Tributary area A (72 mi<sup>2</sup>), which extends from the Sangre de Cristo Range divide to the basin-fill boundary, receives the most precipitation within the study area (fig. 3) and is consequently heavily vegetated. Tributary area B (81 mi<sup>2</sup>), on the other hand, generally extends from the Wet Mountain divide to the basinfill boundary and is characteristically dry and sparsely vegetated except at higher altitudes. Runoff from the two tributary areas was determined using the estimated 1931-60 mean annual runoff for the six gaging stations (table 4), with consideration given to the relative characteristics of these basins and the ungaged area. Runoff data from several other hydrologically similar gaged basins outside the Wet Mountain Valley also were considered. The average annual yield during 1931-60 was estimated to be 50,000 acre-ft from tributary area A and 4,000 acre-ft from tributary area B.

Surface-water outflow from both the entire drainage basin and the basinfill part of the basin was measured by the Grape Creek near Westcliffe gaging station (07095000). During the 1931 through 1960 calendar years, the flow past this gage averaged about 22,000 acre-ft annually.

#### Ground-Water Underflow

Ground-water underflows into or out of both water-budget areas are assumed negligible because low-permeable pre-Tertiary rocks are located along most of the perimeter of the basin. The potentiometric map (pl. 1) indicates that ground-water does not flow across the Grape Creek-Texas Creek drainage divide or the Promontory Divide (Custer-Huerfano County line). Consequently, there are no significant ground-water underflows out of the water-budget areas.

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		Water in				Water out		
urape ureek budget area	Source	Acre-feet	Inches	Percent- tage	Source	Acre-feet	Inches	Percent- age
Entire drainage area <sup>1</sup>	Precipitation <sup>2</sup> Surface water Ground-water underflow <sup>4</sup>	331,000 0	19.4 0 0	0 0 0	Evapotranspiration <sup>3</sup> Surface water <sup>2</sup> Ground-water underflow <sup>4</sup>	309,000 22,000	18.1 1.3	93 7 0
	-		-		Change in storage <sup>4</sup>	0	0	0
	lotals	331,000	19.4	100	lotais	331,000	19.4	100
Basin-fill area <sup>5</sup>	Precipitation <sup>2</sup> Surface water	151,000	17.0	74	Evapotranspiration <sup>3</sup>	183,000	20.5	89
	(tributary area A) <sup>6</sup> Surface-water	50,000	5.6	24	Surface water <sup>2</sup>	22,000	2.5	11
	(tributary area B) <sup>6</sup>	4,000	0.4	2	Ground-water underflow <sup>4</sup>	0	0	0
	underflow <sup>4</sup>	0	0	0	Change in storage <sup>4</sup>	0	0	ο
	Totals	205,000	23.0	100	Totals	205,000	23.0	100
<sup>1</sup> Drainage ba 320 mi2	isin above Grape Cre	eek near We	stcliffe	0005600)	), location shown on fig	gure 11. TI	nis is a	n area of
<sup>2</sup> Based on re <sup>3</sup> Ertimated a	scorded data.		atad for		100			

<sup>5</sup>Estimated as all water otherwise unaccounted for in the budget. <sup>4</sup>Assumed negligible. <sup>5</sup>That part of the Grape Creek near Westcliffe (07095000) drainage composed of basin-fill material (see figure 11). This is an area of 167 mi<sup>2</sup>. <sup>6</sup>Estimated (see text).



Figure 10.-- Locations of water-budget tributary areas and area of basin-fill material in the Grape Creek basin.

#### Evapotranspiration

The evapotranspiration component of the water budgets was estimated to be all water unaccounted for in the budget. The result of this determination is 309,000 acre-ft or 18.1 in. for the entire drainage basin, and 183,000 acreft or 20.5 in. for the basin-fill area. These estimates are in close agreement with the results of several detailed studies conducted elsewhere in Colorado. Wymore (1974) estimated that the average annual evapotranspiration from the 142-mi<sup>2</sup> upper Piceance Creek basin is 18.47 in. Kruse and Haise (1974) measured 23.2 to 24.5 in. of evpotranspiration from high mountain meadows in the South Park area, and 23.5 to 27.8 in. in the Gunnison area.

#### Change in Storage

Because there are no large surface-water reservoirs in the water-budget area and ground-water levels measured in wells during the 24-month period were relatively stable, this component of the budgets was assumed to be zero.

#### WATER QUALITY

#### Ground Water

Water samples were collected and analyzed from 20 wells within the basin (pl. 1). The results of these analyses are shown in table 7. (The system of numbering water-quality sampling sites is explained in the Supplemental Information section of this report.) The analyses indicate that, except for four wells discussed below, the constituents found in ground water in the basin are within the recommended limits for drinking water (table 8) set by the U.S. Public Health Service (1962) and the U.S. Environmental Protection Agency (1972).

The concentrations of dissolved solids, 703 mg/L (milligrams per liter), and dissolved nitrite plus nitrate, 16 mg/L, exceeded the recommended limits of 10 mg/L in water from well SC 22-72-31ADB. As this well is a 15-ft deep stock well, located in a feed lot, the water chemistry is not typical of the water from the major part of the valley. Dissolved solids greater than 500 mg/L may impart a bitter taste to the water. Dissolved nitrite plus nitrate greater than 10 mg/L is generally indicative of pollution of the water from animal wastes and (or) fertilizers and may cause methemoglobinemia (bluebaby disease) in infants (less than 6 months of age) who drink the water or who are breast-fed by mothers who drink the water.

The main significance of the above sampled well shows the relative ease with which surface and near-surface sources may pollute a normally potable domestic water supply. Therefore, it is important that proper consideration be given in the future to waste-disposal practices and their consequent effects upon the ground-water supply, especially in the 46-mi<sup>2</sup> area in which the depth to water is less than 10 ft.

Table 7. -- Chemical analyses of water from selected wells

[Major aquifer codes: 02, Basin fill; E7, Dry Union Formation; E8, Santa Fe Formation or Group; E9, Tertiary igneous, undivided]

	Local well number	Station number	Date of sample	Depth of well (ft)	Major aqui- fer	Dis- solved silica (SiO <sub>2</sub> ) (mg/L)	Dis- solved alumi- num (Al) (µg/L)	Dis- solved iron (Fe) (µg/L)	Dis- solved manga- nese (Mn) (µg/L)	Dis- solved calcium (Ca) (mg/L)
SC	21-73- 8BDD	381438105335900	72-08-22	66	E8	24	10	30	40	82
S	: 21-73-28BDC	381155105330500	72-08-18	117	Е8	18	0	110	10	34
Š	C 21-73-34DBB	381059105314600	72-08-18	202	н В Ш	17	10	40	0	27
S	C 21-73-35CDB	381042105305800	72-08-22	124	Е8	20	10	20	0	34
Š	C 22-72-14CDA	380752105241500	72-08-16	50	02	29	10	20	0	60
S	C 22-72-17CAC	380805105273800	72-08-16	120	E9	29	10	120	10	43
S	C 22-72-31ADB	380608105282300	72-08-18	15	02	21	20	160	20	47
S	C 22-72-32AAB	380605105271000	72-08-16	89	Е8	27	10	6,000	m	77
S	C 22-73- 6ACB	381017105350800	72-08-22	91	02	15	10	60	30	32
S	C 22-73-12DDC	380847105292500	72-08-18	45	02	22	10	100	20	31
S	c 22-73-36DCC	380519105294600	72-08-16	103	E 8	11	0	240	40	22
S	C 23-72- 6AAB	380513105281400	72-08-18	1,200	н В Ш	25	10	40	0	26
S	C 23-72-20CBC	380201105280600	72-08-16	23	02	14	0	20	0	46
S	C 23-72-23DDD	380142105235000	72-08-16	180	02	31	0	80	10	55
S	C 23-72-24DBC	380155105231500	72-08-16	96	02	29	10	20	0	44
S	C 23-72-26CBB	380114105245800	72-08-16	60	02	22	0	60	0	30
S	C 23-73- 4CAC	380443105332200	72-08-18	95	02	15	10	50	0	41
S	C 24-71- 6BAA	385947105222500	72-08-16	29	ΕŢ	41	10	70	10	100
z	A 45-12- 2DAA	381039105362400	72-08-22	110	ł	19	0	4,200	230	43
Z	A 46-12-12DBC	381548105355800	72-08-22	184	Е8	17	10	20	0	26

	s- ved ho- os- rus P) /L)	01 01 02 02	03 00 01 01	00 01 02 07	03 01 01 03 03 03 03 03 03 03 03 03 03 03 03 03
	Di sol'sol pho pho ((	0			
i	Dis- solved nitrite plus nitrate (N) (mg/L)	0.51 .52 .43 .16 3.6	1.1 16 .00 .12 .11	.06 .09 .38 .38	.42 .07 .01 .01
	Dis- bis- fluo- ride (F) (mg/L)	0.5 	०७५७७	- 2 2 <u>0</u> 9	₩₩. ₩₩.
	Dis- solved chlo- ride (Cl) (mg/L)	7.1 2.3 1.0 25	8.2 69 6.8 1.2	3.86 3.66 3.86 3.86 3.86 3.86 3.86 3.86	3.3 31.9 2.0
	Dis- solved sul- fate (SO <sub>4</sub> ) (mg/L)	82 9.5 21 48	21 63 13 8.5 11	8.5 4.3 12.2 11	7.4 4.7 35 23 2.6
6 6	Alka- linity as CaCO <sub>3</sub> (mg/L)	220 124 93 152	150 294 273 112	70 103 163 185	101 139 214 87
	Car- bonate (CO <sub>3</sub> ) (mg/L)	00000	00000	00000	00000
	Bicar- bonate (HCO <sub>3</sub> ) (mg/L)	263 151 113 149 192	183 359 333 145	85 125 198 225 191	123 170 261 226 106
	Dis- solved potas- sium (K) (mg/L)	2.1 1.4 1.1 2.8	2.6 200 4.8 1.4	1.1 3.7 5.5	1.2 6.1 2.3 7.7
	Dis- solved sodium (Na) (mg/L)	21 7.5 5.6 8.0 12	17 39 7.9 4.4	3.8 15.0 13.0	6.5 4.7 4.6 4.6
	Dis- solved magne- sium (Mg) (mg/L)	14 7.2 7.9 10	8.6 14 5.2 7.3	4.7 8.5 6.5 6.5	4.9 12.8 2.9
	Site num- ber <sup>1</sup>	10 116 22	26 77 70 70 70 70 70 70 70 70 70 70 70 70	66 74 101 108 112	114 118 127 144 153

Table 7.--Chemical analyses of water from selected wells--Continued

Table 7.--Chemical analyses of water from selected wells--Continued

## Table 8.--Recommended limits of selected substances in water supplies

[Adapted from U.S. Public Health Service, 1962. The following chemical substances should not be present in a water supply in excess of the listed concentrations where, in the judgment of the reporting agency and the certifying authority, other more suitable supplies are or can be made available]

Concentration, in			
Milligrams per liter	Micrograms per liter		
250			
500			
1.3			
.30	300		
.05	50		
10			
250			
	Concentr Milligrams per liter 250 500 1.3 .30 .05 10 250		

<sup>1</sup>The recommended control limit for fluoride concentrations is variable and dependent upon the annual average of maximum daily air temperature.

<sup>2</sup>In areas in which the nitrate concentration of water is known to be in excess of the listed concentration, the public should be warned of the potential danger of using the water for infant feeding.

The concentration of dissolved iron,  $6,000 \ \mu g/L$  (micrograms per liter), exceeded the recommended limit in water from well SC 22-72-32AAB. Dissolved iron greater than  $300 \ \mu g/L$  imparts a bitter taste to beverages, such as coffee and tea made using the water, and stains laundry and porcelain fixtures.

The concentration of dissolved nitrite plus nitrate of 14 mg/L exceeds the recommended limit in water from well SC 24-71-06BAA,a 29-ft deep windmill.

The concentrations of dissolved iron, 4,200  $\mu$ g/L; manganese, 230  $\mu$ g/L; and fluoride, 1.8 mg/L, exceeded recommended limits in water from well NA 45-12-02DAA. Dissolved manganese greater than 50  $\mu$ g/L causes the same effects as dissolved iron greater than 300  $\mu$ g/L. Dissolved fluoride greater than 1.3 mg/L can cause mottling of children's teeth.

Ground water in the basin is generally hard (as defined by Hem, 1970, p. 225). The hardness, expressed as calcium carbonate (table 7), in the samples analyzed ranged from 74 to 300 mg/L. All but four of the samples have a hardness of 100 mg/L or greater. According to Hem (1970), the hardness of water to be used for domestic purposes becomes objectionable when the concentration is about 100 mg/L.

According to J. M. Klein, P. A. Emery, L. A. Hershey, and R. D. Penley (written commun., 1977), there are two different water-quality zones in the basin-fill aquifer. Generally, east of the axis of the valley the concentration of dissolved silica is greater than 25 mg/L and the concentrations of dissolved solids are greater than 200 mg/L, while in the western part of the basin these constituents are generally less than these concentrations. The difference in water quality between these two zones is probably due to the different sources of the sediments in the two areas. The source of the sediments in the western one-half of the valley is the Sangre de Cristo Range and the source for the east side of the basin is the Wet Mountains.

Dissolved silica in ground water originates from dissolution of various silicate minerals or volcanic glass which the ground water contacts. Large areas of the Wet Mountains are covered with layers of volcanic rocks (fig. 2) that contain substantial proportions of readily dissolved silica. Sediments derived from these volcanic rocks are the probable source of the larger silica concentrations on the east side of the valley.

#### Surface Water

Two surface-water samples were collected for water-quality analyses from a site along lower Grape Creek (pl. 1). The results of these analyses are listed in table 9 and indicate that, at the time the samples were collected, the substances analyzed for were within the recommended limits for drinking water set by the U.S. Public Health Service (1962) (table 8). Wentz (1974) lists temperature, pH, and specific-conductance data for several surface-water sites in the study area.

#### Suitability of Water for Irrigation

The suitability of water for irrigation is dependent on several variables, the most important of which are the quality of the water, the type of soil and crops to be irrigated, and the irrigation practices used. Α concentration of dissolved solids greater than 1,000 mg/L (Robinove and others, 1958) may affect crop growth indirectly through the effect on the physical condition of the soil, and directly by increasing the osmotic pressure of the soil solution and by an accumulation of toxic quantities of various ions within the plant. Thus, the ability of a plant to use water is diminished and a reduced crop yield results (Reeve, 1957, p. 175). Sodium buildup in the soil may reduce soil permeability, which limits air penetration to the root zone, leads to waterlogging, and reduces plant growth (Klein and Bingham, 1975).

A diagram used for classification of irrigation water, based on the sodium-adsorption ratio (SAR) and the specific conductance, is shown on figure 11. This diagram is a useful tool in predicting potential sodium and salinity hazards involved in the use of water for irrigation. The diagram has 16 different categories for classifying irrigation water. These categories range from low salinity (Cl), low sodium (Sl), which is water that can be

	Dis- solved calcium (Ca) (mg/L)	41	42	Dis- solved ortho- phos- phorus (P) (mg/L)	0.01 .02	Carbon diox- ide (CO <sub>2</sub> ) (mg/L)		
	Dis- solved manga- nese (Mn) (µg/L)	20	10	Dis- olved trite plus trate (N) mg/L)	0.09 .16	Tem- pera- ture (de- grees Cel- sius)	22.0 .5	
	Dis- solved iron (Fe) (µg/L)	20	50	s- ved ni uo- ni de ni (/L) (	4	Hď	7.9 7.1	
	Dis- solved alumi- num (A1) (µg/L)	;	1	iis- Di lved sol tho- fl ide ri ide ri (F) (F C(1) (F	3.3 0. 3.1 .	Speci- fic con- ductance (µmhos/cm at 25°C)	291 286	
	Dis- solvec silica (siO <sub>2</sub> ) (mg/L)	22	21	Dis- D olved sc sul- c fate r (SO <sub>4</sub> ) ( mg/L) (n	8.0 12	Sodium- ad- sorp- tion ratio (SAR)	0.4 .4	
and the second se	Dis- charge (ft <sup>3</sup> /s)	4.0	10	lka- nity <sup>s</sup> as g/L) (	144 149	Per- cent so- dium	14 13	
	Date of sample	.2-08-18	2-11-30	ar-A nate li 03) C ng/L) (m	0 0	Non- car- bonate hard- ness (mg/L)	00	
	umber	262000 7	262000 7	3icar- C oonate bc (HCO <sub>3</sub> ) (C (mg/L) (m	176 182	Hard- ness as CaCO <sub>3</sub> (Ca,Mg) (mg/L)	130 140	
	Station r	3807521052	3807511052	Dis- solved E potas- t sium (K) (mg/L)	2.2 1.6	Dis- solved solids (tons per acre-ft)	0.25 .25	
	n name	reek at t Lake	reek at t Lake 	Dis- solved sodium (Na) (mg/L)	-10 9.5	Dis- solved solids (sum of constit- uents) (mg/L)	182 187	ure 2.
and the second	Statio	Grape C Hermi Road-	Grape C Hermi Road-	Dis- solved magne- sium (Mg) (mg/L)	7.7 8.0	Dis- solved ortho- phos- phate (PO4) (mg/L)	0.01 .06	l See fig
	Site num- ber <sup>l</sup>	-	-	Site num- ber <sup>1</sup>		Site num- ber <sup>1</sup>		

Table 9. --Chemical analyses of water from selected surface-water sites



Figure 11.-- Classification of irrigation water.

applied to nearly all types of soils and crops with little danger of harmful effects, to very high salinity (C4), very high sodium (S4), which is water suitable for irrigation use only under very limited conditions (table 10).

All of the water samples analyzed during the course of this study plot within the Sl sodium-hazard category and all but one plot within either the Cl or C2 salinity-hazard categories (fig. 11). This would indicate that both the surface and ground water have low-sodium hazard and low- to medium-salinity hazard. Within the study area, surface and ground water are generally safe to use for irrigation under most conditions.

#### SUMMARY

The Wet Mountain Valley is an intermontane trough filled with unconsolidated deposits of Cenozoic age, and bordered by mountain ranges composed of volcanic rocks of Tertiary age and igneous, metamorphic, and sedimentary rocks of pre-Tertiary age. The valley includes an area of 540 mi<sup>2</sup> and is primarily drained by Grape Creek and Texas Creek. Average annual precipitation ranges from less than 16 in. to more than 40 in., generally increasing with land-surface altitude.

Ground water occurs in two different hydrologic areas within the study area--the mountainous area, which is composed chiefly of consolidated rock; and the valley-floor area, which is underlain by unconsolidated basin-fill sediments. In the mountainous area, well yields are highly unpredictable but are generally less than 10 gal/min. In the basin-fill aquifer, yields of greater than 50 gal/min can generally be expected from wells located near the central part of the valley. In areas around the edge of the basin-fill aquifer, well yields of less than 50 gal/min are generally reported.

The basin-fill deposits cover an area of about 233 mi<sup>2</sup> and are the principal source of ground water in the study area. South of Westcliffe, this material is at least 6,700 ft thick and the upper 1,200 ft is composed of sand and gravel with interbedded and intergranular clay. There is an estimated 1.5 million acre-ft of ground water stored in the upper 100 ft of saturated basin-fill sediments although considerably less than this amount could be developed from wells. Hydrographs of selected wells indicate that the volume of water stored in the basin-fill material is not changing.

Ground water occurs under both artesian and water-table conditions within the basin-fill aquifer. Artesian conditions exist, to some extent, throughout most of the basin-fill area and there are several flowing wells. The depth to the water table is less than 10 ft in an area of about 46 mi<sup>2</sup> along the central part of the valley and is less than 100 ft in most of the remainder of the valley.

Ground-water movement in the basin-fill aquifer is generally toward Grape and Texas Creeks. Ground water leaves the aquifer either by discharge to streams, discharge by wells, or as direct evapotranspiration from the water table. Underflow out of the basin is considered negligible.

COLORADO WATER RESOURCES LIBRARY U.S. GEOLOGICAL SURVEY WRD, DEN Table 10.--Salinity and sodium classification for irrigation water

<del></del>		SALI	ΝΙΤΥ	
	Class	Specific conductance, in micromhos per centimeter at 25°C	Comments	
C1	(low)	Less than 250	May be used for any crop on nearly all soils.	
C2	(medium)	250-750	May be used if a moderate amount of leaching occurs.	
C3	(high)	750-2,250	Should be used only on soils with bet- ter than adequate drainage and for plants with good salt tolerance.	
С4	(very high)-	Greater than 2,250	Not suitable for irrigation under ordi- nary conditions, but may be used on permeable soil with adequate drainage if water is applied in excess to pro- vide considerable leaching. Very salt-tolerant plants should be selec- ted.	
		SOD	IUM	
	Class Comments			
<b>S</b> 1	(low)	May be used for any crop on nearly all soils.		
S2	(medium)	May be used on coarse-textured (sandy) soils with good per- meability; may present a moderate sodium problem in fine- textured (clay) soils unless gypsum is present.		
S3	(high)	May present sodium pr	oblems in most soil types.	
S4	(very high)-	Not suitable for irri under special condi	gation of any crops in any soil, except tions.	
		Sodium-adsorption-rat specific conductanc	io range for each class is dependent on e.	

Streamflow in the area is typical of a high mountain valley with most runoff occurring in June and July as the result of snowmelt. Observed streamflow consists of runoff from precipitation, return flow from irrigation, discharge of ground water, and water released from reservoir storage. Continuous records of streamflow have been collected since 1924 at the Grape Creek near Westcliffe gaging station. Additionally, two continuous-record and four partial-record stations have been operated in the study area since August 1974. Analysis of the streamflow data for water years 1975 and 1976 shows that the average annual runoff for 1975-76 at the seven stations ranged from about 420 to 24,240 acre-ft per year. Long-term annual runoff (1931-60) at these stations ranges from an estimated  $0.02 \, (\text{ft}^3/\text{s})/\text{mi}^2$  to an estimated 1.17  $(ft^3/s)/mi^2$ , generally increasing with station altitude. Using the data at two partial-record stations where natural-flow conditions prevailed, the magnitude, frequency, and depth of floods were estimated using basin characteristics and regionalized regression equations.

DeWeese Reservoir, located near the study area, has a storage capacity of 1,772 acre-ft. All of the other reservoirs within the study area have a combined storage capacity of about 215 acre-ft.

Generalized annual water budgets for two areas in the Wet Mountain Valley indicate that surface-water outflow is only 7 to 11 percent of the total water supply from precipitation and other sources. The remaining water is lost to the atmosphere by evapotranspiration. Ground-water underflow and changes in ground-water storage are considered insignificant in both water budgets.

The quality of both the surface and ground water is generally within the recommended limits for drinking water set by the U.S. Public Health Service (1962). The hardness of the ground water is generally greater than 100 mg/L as calcium carbonate. There are two different water-quality zones in the basin-fill aquifer. East of the axis of the valley, the concentration of dissolved silica is greater than 25 mg/L and the concentration of dissolved solids is greater than 200 mg/L. In the western part of the valley, these constituents are generally less than these concentrations.

Both the surface and ground water are generally safe to use for irrigation. None of the samples analyzed indicated a potential sodium hazard and most indicated only a low- to medium-potential salinity hazard.

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#### SUPPLEMENTAL INFORMATION

#### System of Numbering Water-Quality Sampling Sites

The system of numbering water-quality sampling sites is based on the grid system of latitude and longitude. The system provides the geographic location of the site and a unique number for each site. The number consists of 15 digits. The first six digits denote the degrees, minutes, and seconds of latitude; the next seven digits denote degrees, minutes, and seconds of longitude; and the last two digits are a sequential number for sites within a l-second grid. (In the event that the latitude-longitude coordinates for two or more sites are the same, sequential numbers "01," "02," and so forth are assigned. See figure 12.)

#### System of Numbering Wells

The well locations in this report are given numbers based on the U.S. Bureau of Land Management system of land subdivision, and show the location of the wells by quadrant, township, range, section, and position within the section. A graphic illustration of this method of well location is shown in figure 13. The first letter of the location number indicates which principal meridian governs the area in which the well is located; S indicates the sixth principal meridian and N indicates the New Mexico principal meridian. The second letter indicates the quadrant in which the well is located. Four quadrants are formed by the intersection of the base line and the principal meridian; A indicates the northeast quadrant, B the northwest, C the southwest, and D the southeast. The first number indicates the township; the second, the range; and the third, the section in which the well is located. The letters following the section number indicate the location of the well within the section. The first letter denotes the quarter section; the second, the quarter-quarter section; and the third, the quarter-quarter-quarter section. The letters are assigned within the section in a counterclockwise direction, beginning with A in the northeast quarter. Letters are assigned within each quarter section and within each quarter-quarter section in the same manner. Where two or more locations are within the smallest subdivision, consecutive numbers, beginning with 1, are added in the order in which the wells are inventoried. For example, SC23-72-15DAA indicates a well in the NE<sup>1</sup>/<sub>4</sub>NE<sup>1</sup>/<sub>5</sub>SE<sup>1</sup>/<sub>4</sub> sec. 15, T. 23 S., R. 72 W.

#### System of Numbering Surface-Water Stations

Surface-water station numbers are assigned in downstream direction along the main stream. Station numbers on tributaries are assigned between station numbers on the main stream in the order in which those tributaries enter the main stream. Downstream-order station numbers are not assigned to miscellaneous sites where water-quality samples are collected or discharge measurements are made intermittently.









Figure 13.--System of numbering wells.

SC 23-72- 6AAB1 (Altitude 7,930 feet above mean sea level)	Thick- ness, in feet	Depth below land surface, in feet
Soil, brown; contains scattered 1-inch rounded gravel	3	3
Gravel, clayey; contains rounded cobbles as large as 6 inches	- 3	6
Gravel, very fine, to very coarse sand; contains some medium gravel	- 4	10
Sand, very coarse, silty; contains some medium gravel	- 10	20
Sand, very coarse, arkosic; contains some reddish-brown clay	- 10	30
Clay, reddish-brown	- 5	35
Gravel, fine to very fine, clayey	- 10	45
Clay, reddish-brown	- 5	50
Gravel, fine to very fine, clayey	- 5	55
Gravel, fine, well-sorted, clayey	- 5	60
Sand, very coarse, clayey, grading to well-sorted fine gravel	- 10	70
Clay, reddish-brown, sandy, grading to clay; contains thin yellowish-gray, hard, platy bentonite(?) stringers	- 10	80
Gravel, very fine, to coarse sand, clayey; contains cemented beds in bottom 10 feet	- 25	105
Gravel, very fine, to sand, very coarse, clayey	- 15	120
Gravel, very fine to fine, clayey	- 40	160
Gravel, grading from coarse to very coarse sand; contains clay streaks and some intergranular clay	- 20	180
Gravel, very fine to fine, and sand, clayey; contains some clean beds	· 30	210
Clay, grayish-yellow, and interbedded sand and gravel	- 20	230
Sand and very fine gravel, some clayey	- 20	250
Sand, very coarse, to gravel, very fine, clayey; contains cemented zone from 255 to 260 feet	- 20	270
Sand, very coarse to coarse, grading to fine gravel	- 10	280
Gravel, very fine to fine, and sand grading to very fine, silty gravel	· 20	300
Sand, silty, poorly sorted	· 25	325
Gravel and sand, clayey	- 10	335

# Geologic Log of Test Hole SC 23-72- 6AAB1

SC 23-72- 6AAB1 (Altitude 7,930 feet above mean sea level)	Thick- ness, in feet	Depth below land surface, in feet
Sand, coarse to medium, clayey; contains cemented(?) sand and gravel	- 40	375
Sand, medium, to very fine gravel, clayey; clean(?) from 375 to 380 feet	- 20	395
Sand, very clayey, grading to clay, sandy	- 20	415
Sand and interbedded clay	- 20	435
Clay, sandy; contains some very fine gravel	- 20	455
Sand, coarse, clayey, grading to clay, containing very fine gravel	- 20	475
Sand, very coarse, grading to sandy clay	- 20	495
Sand, medium, to very fine gravel, grading to sandy clay	- 20	515
Sand, medium to coarse, grading to sandy clay	- 20	535
Clay, sandy, and interbedded clay	- 20	555
Sand, medium to coarse; contains clay streaks	- 20	575
Sand, poorly sorted, some clayey and reddish-brown clay streaks; some caliche bottom 5 feet	- 40	615
Sand, medium to very coarse, and very fine gravel; contains clay streaks	- 20	635
Sand, medium to coarse, clayey; contains beds of yellowish- gray bentonitic(?) clay	- 20	655
Sand, very coarse, clean, grading to poorly sorted clayey sand	- 20	675
Sand, coarse, to very fine gravel, clayey	- 50	725
<pre>Sand, silty to clayey, cemented(?); small amount of reddish- brown clay and yellowish-gray, bentonitic(?) clay</pre>	- 10	735
Sand, silty to clayey, loosely cemented, grading to reddish- brown clay	- 15	750
Clay, reddish-brown to greenish-yellow	- 10	760
Clay, sandy, greenish-yellow to pink; contains some caliche	- 15	775
Clay, reddish-brown, sandy, grading to varicolored clays	- 20	795
Sand and very fine gravel, clay-filled	- 20	815
Sand, medium to coarse, clayey, grading to sandy clay	- 20	835

# Geologic Log of Test Hole SC 23-72- 6AAB1--Continued

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SC 23-72- 6AAB1 (Altitude 7,930 feet above mean sea level)	Thick- ness, in feet	Depth below land surface, in feet
Sand, fine to medium; contains some clay grading to very fine gravel and coarse sand	- 20	855
Sand, fine, clayey, grading to clean coarse sand	- 20	875
Sand, medium to coarse, grading to clayey sand and some tightly cemented conglomerate	- 20	895
Sand and gravel, conglomeratic; some interbedded clay	- 20	915
Sand to very fine gravel; contains loosely cemented stringers-	- 20	935
Sand, clayey; contains some tightly cemented sand and gravel beds	- 20	955
Sand, poorly sorted to clayey	- 20	975
Sand, medium to very coarse, less clayey than above	- 20	995
Sand, very coarse, grading to very fine, clayey gravel	- 20	1,015
Sand and fine gravel grading to sand and medium gravel; contains clay stringers	- 20	1,035
Sand and very fine gravel grading to tightly cemented gravel	- 20	1,055
Sand, fine to very coarse; contains some reddish-brown clay and cemented beds	- 20	1,075
Gravel, very fine to medium; contains cemented beds and sandy clay	- 40	1,115
Gravel, fine to very fine, and sand, and interbedded caliche, brown clay, and sandstone	- 20	1,135
Sand, fine to very coarse, silty and clayey	- 20	1,155
Sand and interbedded reddish-brown clay	- 25	1,180
Sand, silty to clayey	- 20	1,200

# Geologic Log of Test Hole SC 23-72- 6AAB1--Continued

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Records of Selected Wells

[Depth to water: F, Flowing. Major aquifer codes: 02, Basin fill; E7, Dry Union Formation; E8, Santa Fe Formation or Group; E9, Tertiary, igneous, undivided; F1, Precambrian, undifferentiated. Type lift: C, Centrifugal; J, Jet; N, None; P, Piston; S, Submersible; T, Turbine. Power: E, Electric; G, Gas, diesel, natural or LP gas engine; H, Hand; N, None; W, Wind.

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Records of Selected Wells--Continued

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Major aqui- fer	02 02 02 E8	E8 F1 02 02	02 02 02 02 02	02 02 12 12 12 12 12 12 12 12 12 12 12 12 12	02 02 02 02 02	E 8 02 02 02 02
Date water level meas- ured (M-Y)	12-64  8-71 4-71 11-70	11-70 4-71 4-71 4-71 4-71 4-71	11-70 11-70 11-70 7-71 5-74	11-70 11-70 11-70 11-70 11-70	8-71 10-70 10-70 10-70 11-70	11-70 12-70 12-70 10-70 12-70
Depth to water (ft)	* <del>+</del> 5 - 5	۳۶ <sup>7</sup> 7 9	₩₩4 ₩ Q	-44 00	m -  -	3 85 112 44
Well depth (ft)	128 89 13 26	103 18 110 300 13	30 16 12 1,200 92	34 40 16 25 180	22 28 115 30	1 5 5 5 4 4 4 5 7 5 4 5 7 2 7 2 7 2 7 5 7 4 7 5 7 5 7 5 7 5 7 5 7 5 7 5 7 5
Alti- tude of land surface (ft)	8,040 8,238 8,240 8,182 7,990	8,020 8,635 8,650 8,400 8,400	7,975 7,905 7,895 7,930	7,995 8,043 8,000 7,995 8,130	8,132 8,142 8,142 8,180 8,175	8,035 8,070 8,000 8,020 8,020
Year drilled	1964 1953 	 1861 1971  1953	1958 1870  1971			1962 1960 1952 
Local well number	SC 22-73-27BAD SC 22-73-28BBC SC 22-73-28BDD SC 22-73-35CDB SC 22-73-36DAA	SC 22-73-36DCC SC 23-71- 4DCC SC 23-71- 9ABC SC 23-71- 9DBB SC 23-72- 1AAB	SC 23-72- 4CDB SC 23-72- 5AAB1 SC 23-72- 5ABB2 SC 23-72- 6AAB1 SC 23-72- 6AAB1	SC 23-72- 6BDD SC 23-72- 6CBD SC 23-72- 7ADA1 SC 23-72- 7ADA1 SC 23-72- 7ADA2 SC 23-72- 7BCC1	SC 23-72- 7BCC2 SC 23-72- 7CBB1 SC 23-72- 7CBB2 SC 23-72- 7CBB2 SC 23-72- 7CCC1 SC 23-72- 7CCC1	SC 23-72- 8CCC SC 23-72-108BA SC 23-72-108DD SC 23-72-10BDD SC 23-72-10DDD
Site number on figure 2	61 62 64 65	66 67 69 70	71 74 75	76 77 78 80 80	81 83 84 85	88 87 89 90 90

<b>III</b> -	~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~	NNDHH	ເທດທາ	NHNNH	ν-τττ	x
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4,50	15	50	16	20	20	14
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12-70 11-70 11-70 11-70 11-70	5-74 10-70 11-70 11-70 11-70	11-70 7-71 7-71 11-70	11-70 12-70 11-70 12-70 11-70	11-70 11-70 9-70 11-70	11-70 11-70 4-71 11-70 8-71	9-70 9-70 8-71 11-70 12-70
8 7 8 7 9 8	ღოთნი	00 <sup>4</sup> - 100	547.95 747.05	~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~	22 22 22	67 68 34 187
35 45 30 245 245	75 22 83 13	23 23 65 40	51 50 180 23 23	160 96 82 60 16	100 100 7 52	113 70 57 8 220
8,048 8,005 7,995 8,025 8,050	8,059 8,082 8,165 8,375 8,154	8,200 8,200 8,198 8,078 8,105	8,061 8,073 8,125 8,215 8,130	8,135 8,181 8,185 8,151 8,200	8,215 8,182 8,800 8,300 8,300	8,895 8,865 8,405 8,455 8,455
   1963	1974  1961 	 1947 1951	1960  1940 1960	1960	1964  1970 1966	1970
SC 23-72-14CAA SC 23-72-15BBB SC 23-72-16ABA SC 23-72-16CAC SC 23-72-16CAC SC 23-72-17BBC	SC 23-72-18AAD SC 23-72-18ABB SC 23-72-19AAD SC 23-72-19BCC SC 23-72-19BCC SC 23-72-20BBC	SC 23-72-20CBC1 SC 23-72-20CBC2 SC 23-72-20CBC3 SC 23-72-20BC8 SC 23-72-21BC8 SC 23-72-21BC8	SC 23-72-22BAC SC 23-72-23BCB SC 23-72-23DD SC 23-72-24ABB SC 23-72-24CC	SC 23-72-24CDB SC 23-72-24DBC SC 23-72-24DDA SC 23-72-26CBB SC 23-72-33DAA1	SC 23-72-33DAA2 SC 23-72-36ADB SC 23-73- 4CAC SC 23-73-13ACC SC 23-73-15AAB	SC 23-73-15ACD SC 23-73-15ADB SC 23-73-23ADC SC 23-73-24ABB SC 24-71- 4CAD
92 94 95	96 99 99 99	101 102 103 104	106 107 108 1109	1124	1116 1117 1119 1120	121 122 123 124

Records of Selected Wells--Continued

Use of water	ຑຑຑຑຑ	νυντυ	νννντ	ר דדי א	~ ~ ~ ~ ~ ~ ~	<b>エエーー</b>
Power	33333	333z3	<u>س د، س د د</u>	ו שים צ	33131	21122
Type lift		0 0 0 Z 0	5 5 C C S	4 4 0 9 म	66F61	ZIIZZ
Casing diam- eter (in.)	36	രവരരാ	موماية	no o o o M	ه و ا م م ا	88800
Draw- down (ft)			135  10	2		
Yield (gal/ min)	10		2 20 10	5	20	40 440 9
Major aqui- fer	E8 E7 E7 E8 E8	E7 E7 E7 E8	Е8 02 E8 E8 E8	02 68 1 - 2 8	02 10 10 10 10 10 10 10 10 10 10 10 10 10	н н н н н н н н н н н н н н н н н н н
Date water level meas- ured (M-Y)	12-70 4-71 4-71 10-70 4-71	4-71 4-71 4-71	6-66  11-70 11-70 7-69	12-59 4-71 11-70 	5-74 5-74 10-70 10-70 10-70	11-70 10-70 10-70 10-70 10-70
Depth to water (ft)	90 25 138 24 219	258 258 320 309	165  92 180	295 Dry 	143 133 51 24 F	և. և. և. և. և.
Well depth (ft)	99 29 142 70 228	 240 332	400  142 150 220	310 387 12 110 150	145 130 196	161 159 184 89 121
Alti- tude of land surface (ft)	8,320 8,240 8,358 8,721 8,480	8,438 8,485 8,465 8,571 8,585	9,038 875 8,320 8,545 8,440	8,495 8,755 8,170 8,835 7,650	7,860 7,880 7,749 7,620	7,630 7,640 7,650 7,600
Year drilled	 1965	  1931	1966  1964 1969	1959  1954		
Local well number	SC 24-71- 5BCA SC 24-71- 6BAA SC 24-71- 7ACD SC 24-71-14DCA SC 24-71-14DCA	SC 24-71-17ABA SC 24-71-17CBA SC 24-71-18DAA SC 24-71-19BDD SC 24-71-22BBC	SC 24-71-25CBA SC 24-71-35BAB SC 24-72- 4DBB SC 24-72- 5ABB SC 24-72- 5ABB	SC 24-72-15AAD SC 24-72-26BAC SC 24-72-28CAA SC 24-72-28CAA NA 45-12- 2DAA NA 46-12- 3DDD	NA 46-12-10BDD NA 46-12-10DCB NA 46-12-11ACC NA 46-12-11CDA NA 46-12-11CDA	NA 46-12-12DBB1 NA 46-12-12DBB2 NA 46-12-12DBB2 NA 46-12-12DBC NA 46-12-12DBD1 NA 46-12-12DBD1
Site number on figure 2	126 127 128 129 130	131 132 133 134 135	136 137 138 138 140	141 142 144 144	146 147 148 149	151 152 153 154

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10-70	4-71	10-70	10-70	5-74	5-74 5-74
LL.	8	24	23	86	35 63
131	23	108	163	1	60 70
7,625	7,740	7,879	7,910	7,935	8,330 8,465
9 9 9 9	1	8	1	1	
NA 46-12-12DDB	NA 46-12-13ACB	NA 46-12-14BBD	NA 46-12-15AAC	NA 46-12-15ABC	NA 46-12-22DDA NA 46-12-27ADB
156	157	158	159	160	161 162

Site number on figure 2	U.S. Geological Survey station number	Station name	Date of sample	Temperature (degrees Celsius)	Specific conductance (µmho/cm at 25°C)
1	07093980	Lake Creek near Hillside, Colo.	7-31-74 8-19-74 9-16-74 10-16-74 11-15-74	11.5 12.5 9 7 3	
			12-18-74 4-18-75 5-21-75 6-13-75 8- 6-75	0 2 11.5 10 10.5	150
			9-9-75 10-7-75 11-11-75 12-23-75 2-3-76	10 8 1.5 1.5 2.5	119 130 123
			4- 8-76 5- 5-76 8-13-76 9-17-76	7 9.5 10.5 9.5	122 127 120 125
2	07094530	Grape Creek near Bradford, Colo.	7-31-74 8-19-74 10-16-74 11-15-74 12-18-74	19 12.5 3 0 0	
			5-20-75 6-13-75 7-10-75 8- 6-75 9- 9-75	9 6.5 9.5 11.5 5.5	101
			10- 7-75 11-11-75 12-23-75 2- 3-76 4- 8-76	9.5 0 0.5 0.5 2.5	155
			5- 5-76 6-30-76 8-13-76 9-17-76	5.5 14 12.5 20	101 115 139 115

# Specific Conductance and Temperature of Water Data fromSelected Surface-Water Stations

Site number on figure 2	U.S. Geological Survey station number	Station name	Date of sample	Temperature (degrees Celsius)	Specific conductance (µmho/cm at 25°C)
4	07094600	South Colony Creek near Westcliffe, Colo.	8- 1-74 8-19-74 9-16-74 10-16-74 11-15-74	14 10.5 6.5 3 0	
			4-17-75 5-20-75 6-13-75 7-10-75 8- 6-75	0.5 4.5 6.5 7.5 9	140
			9- 9-75 10- 7-75 11-11-75 12-23-75 2- 3-76	9 6.5 0 0.5	95 100 122
			4- 8-76 5- 5-76 6-11-76 7-12-76 8-13-76	2 3.5 8 13.5 8.5	116 91 70 70 90
			9-17-76	10.5	105
5	07094700	Antelope Creek near Rosita, Colo.	7-31-74 4-17-75 5-20-75 6-13-75 7-10-75	30.5 10.5 9 12 14.5	310
			8- 6-75 9- 9-75 11-11-75 2- 3-76 4- 8-76	16 12 0 0 9.5	360 362
			5- 5-76 6-30-76	10 23	370 370

# Specific Conductance and Temperature of Water Data from Selected Surface-Water Stations--Continued

Site number on figure 2	U.S. Geological Survey station number	Station name	Date of sample	Temperature (degrees Celsius)	Specific conductance (µmho/cm at 25°C)
6	07094800	Venable Creek near Westcliffe, Colo.	7-31-74 8-19-74 9-16-74 10-16-74 5-21-75	10.5 9.5 5.5 3.5 6	
			6-13-75 7-10-75 8- 6-75 9- 9-75 10- 7-75	8 9 10 9 6	138 70
			11-11-75 12-23-75 2- 3-76 4- 8-76 5- 5-76	0 0 1 6	170 74
			6-11-76 6-30-76 8-13-76 9-17-76	7 11 9.5 9	110 120 140 160
7	07094900	Middle Taylor Creek near Westcliffe, Colo.	8- 1-74 8-19-74 9-16-74 10-16-74 11-26-74	16 15.5 7.5 9 0	240
			1-29-75 6-13-75 8- 6-75 9- 9-75 10- 7-75	0 13 17.5 13.5 10.0	197 215
			10-16-75 12-23-75 2- 3-76 4- 8-76	7 0 0.5	218 220 215
			5- 5-76 5-11-76 6-11-76 6-30-76 7-12-76 8-13-76	6 7.5 10.5 17 12 13	150 125 105 155 50 185
			9-17-76	10	210

# Specific Conductance and Temperature of Water Data from Selected Surface-Water Stations--Continued

UNITED STATES DEPARTMENT OF THE INTERIOR GEOLOGICAL SURVEY



MAP SHOWING LOCATION OF WELLS, SURFACE-WATER GAGING STATIONS, AND POTENTIOMETRIC SURFACE OF THE BASIN-FILL AQUIFER, WET MOUNTAIN VALLEY, PARTS OF CUSTER AND FREMONT COUNTIES, COLORADO