GEOHERMAL DEVELOPMENT OF THE SALTON TROUGH, CALIFORNIA AND MEXICO

Editors:

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GEOHERMAL DEVELOPMENT OF THE SALTON TROUGH, CALIFORNIA AND MEXICO

Abstract

A geological description is given of the Salton Trough followed by a chronological history of attempts to exploit the area's geothermal resources. In addition,

detailed descriptions are given of all ongoing geothermal projects in the area and the organizations conducting them.

Introduction*

This report has been prepared for use by the Second United Nations Geothermal Symposium to be held in San Francisco, California, 20-29 May 1975. The University of California, in association with other public and private agencies, is supporting the United Nations Secretariat in the organization and presentation of this symposium. The goal of the meeting is to foster the exchange of information and experience among countries involved in geothermal research. Toward this end, this report presents information on geothermal developments in the Salton Trough, California and Mexico (Figs. 1 and 2), one of North America's largest geothermal resource areas. The Salton Trough now has one operating geothermal power plant at Cerro Prieto, Mexico, and intense research and development efforts are underway to establish additional plants.

This paper discusses these activities, including the problems and progress of past and present research projects, and presents a geological description of the region.

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Regional Geology of the Salton Trough†

The water-dominated geothermal fields in the Imperial and Mexicali Valleys of Southern California, U. S. A., and Northern

Baja California, Mexico, occur within the physiographic province known as the Salton Basin. This basin, which forms the

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Fig. 1. Salton Sea and Salton Trough as seen from orbiting Gemini V spacecraft, August 1965. Coachella Valley upper left, Imperial and Mexicali Valleys lower right. (Photo courtesy of NASA.)

Fig. 2. Salton Trough geothermal province and known geothermal resource areas (~400 mi² KGRA).

northern part of the Colorado River delta (Fig. 3), is the surface expression of a deep, sediment-filled, structural trough, or rift valley, called the Salton Trough. It is the landward extension of the Gulf of California northwards into North America. The overall structure of the trough is apparently controlled by the numerous strike-slip faults related to the San Andreas fault system (Fig. 4).

The Gulf of California and the Salton Trough are areas of rapid tectonic deformation, where patterns of high heat flow and seismicity, together with patterns of sedimentation and volcanicity, reflect a transition from the divergent plate boundary of the East Pacific Rise to the transform

boundary represented by the San Andreas fault system (Wilson, 1965; Larson *et al.*, 1968; Atwater, 1970; Elders, *et al.*, 1972; and Moore, 1973).

Paleontological evidence suggests that the Gulf of California has existed as a morphotectonic depression for the last 15 million years. Marine foraminifera of late Miocene to early Pliocene age attest to the existence of deep water in the Gulf from 11 to 8 million years ago (Ingle, in Elders and Biehler, 1975). Marine geophysical surveys have revealed numerous elongate topographic depressions in the Gulf floor (Fig. 4) bounded by seismically active faults. These closed basins have heat flows several times the crustal

Fig. 3. Salton Trough index map showing Imperial and Mexicali Valleys.

average, have positive gravity anomalies of up to +80 mgal, and appear to have been produced by active spreading between en echelon fault segments (Moore, 1973). At the mouth of the Gulf, south of the Tamayo fracture zone, spreading on the East Pacific Rise has proceeded at 6 cm a year for the last 4.5 million years (Larson, 1972). The development of the

deep basins also appears to date from that time.

STRUCTURE

In gross structure and size, the Salton Trough is similar to the deep, closed, marine basins found in the Gulf of California. However, it is partially filled

with a vast accumulation of mainly continental sedimentary rocks. It is a complex rift valley bordered by mountains consisting of Mesozoic, and older, granitic and metamorphic rocks, with some Tertiary volcanic rocks. It has steep, step-faulted margins and a broad, relatively flat basement floor beneath a cover of sedimentary rocks 6 to 7 km thick in the center of the Imperial Valley (Elders, et al., 1972). These rocks are transected by three major fault systems, which trend northwest-southeast: the San Andreas, San Jacinto, and Elsinore fault zones (Fig. 3). Numerous subsidiary blocks and basins are aligned along these major strike-slip faults.

Seismic activity along these faults makes the Salton Trough one of the most earthquake-prone areas in North America. There have been more than 12 earthquakes of magnitude greater than 6.0 (Richter scale) in the area this century. Early in 1975, an earthquake swarm, centered

Fig. 4. Gross tectonic environment of the Salton Trough. The Pacific Coast of North America is dominated by transform fault systems, which connect the spreading centers of the East Pacific Rise to those of the Gorda Ridge. Also shown are pull-apart basins between en echelon fault segments in the Gulf of California. Oceanic fracture zones (FZ) and continental faults (F) are solid black lines, dashed where uncertain. Other abbreviations: SAF = San Andreas Fault; EF = Elsinore Fault; SJF = San Jacinto Fault; ABF = Aqua Blanca Fault; SRF = Santa Rosalia Fault; W = Wagner Basin; D = Delfin Basin;

▲ = Holocene volcanoes; B = Salton Buttes; C = Cerro Prieto; and R = Revillagigedo. (Source: Elders, et al., 1972.)

southeast of Brawley, produced more than 1,000 small earthquakes in only a few days (36 were of Richter magnitude greater than 3).

All of the hypocenters observed in the Salton Trough and the Northern Gulf of California are shallower than 15 km and most originate at less than 6 km. Therefore, nearly all strain release is in the upper 10 km. Work on surface seismic-wave dispersion suggests that the crust beneath the Imperial Valley thins to about 20 km and that the crust under the northern half of the Gulf may be only 10 km thick (Thatcher, in Elders and Biehler, 1975). Certainly, a large part of the regional gravity anomaly can be explained in this way. A complete Bouguer gravity-anomaly map of the Salton Trough indicates that, although the basin is underlain by low-density sediments, it is characterized by a broad gravity maximum, another indication of thin crust (Biehler, *et al.*, 1964 and Elders, *et al.*, 1972).

SEDIMENTARY HISTORY

The present apex of the Colorado River delta forms a low divide (11 m above sea level at its lowest point) between the Imperial Valley to the north and the Mexicali Valley to the south (Fig. 3). Most of the Imperial Valley lies below sea level. At its northern end is the Salton Sea, which covers about 930 km² and has a surface elevation of about 71 m below sea level. Water entering this basin can only escape by evaporation. The Colorado River enters the Salton Trough from the east at Yuma 43 m above sea level. The delta slopes northward (at 0.8 m/km) into the Salton Basin and southward (at

~0.35 m/km) to the Gulf of California (Thompson, 1968). During 1905 to 1907, the Colorado flooded over the delta crest into the Salton Basin, forming the Salton Sea (Sykes, 1937). Although now the natural discharge of the River is into the Gulf of California, inflow of Colorado River water via irrigation canals permits the Salton Sea to persist.

These observations are the key to understanding the history of sedimentation in the basin. The Salton Trough is an actively growing rift valley, in which sedimentation has almost kept pace with tectonism. Formation of the delta perpendicular to the length of the Gulf of California rift has isolated the Salton Basin from the Gulf, forming a closed sedimentary basin 200 km long and up to 130 km wide. Since its formation, the Salton Basin has undergone cycles of filling with freshwater lakes and desiccation as the Colorado River changed course, alternately flowing north or south. Sediments from the walls of the Basin form marginal alluvial fans, but the Colorado River has dominated its sedimentary history. The deltaic deposits consist of interbedded sand, silts, clays, and pebble conglomerates (Van De Kamp, 1973). These rocks are interspersed with lake sediments and reworked eolian deposits. The percentage of sandbodies in the delta sediments decreases away from the delta apex towards the northwest (Randall, 1971). However, rather little is known of the nature and age of the sedimentary rocks in the central part of the Basin. The deepest well yet drilled in the Imperial Valley (Standard Oil of California, Wilson #1) penetrated 4 km of Pleistocene fluvial and lacustrine sediments (Merriam and Bandy, 1965; Muffler and Doe, 1968).

The stratigraphy of the Neogene rocks cropping out in structurally complex zones on both sides of the Imperial Valley have been summarized by Dibblee (1954). Included in these formations are a few marine units, the oldest of which may be as old as Miocene. Maximum marine submergence occurred during the Pliocene, and intermittent shallow marine environments persisted in the western part of the Imperial Valley until middle Pleistocene (Woodward, 1974). Such marine rocks have not been reported from any of the numerous drill holes in the main part of the valley.

THERMAL ANOMALIES

The Salton Trough is characterized by numerous areas of very high temperature gradients at shallow depth. In addition to the Cerro Prieto geothermal field (C in Fig. 4), 8 or 10 anomalously hot zones are known in the Imperial Valley (Fig. 5). The high thermal gradients are thought to be related to circulation of convecting hot groundwater in the thick sedimentary fill (Dutcher, *et al.*, 1972). These thermal anomalies coincide with low-amplitude, positive, residual-gravity anomalies with closures of 2 to 20 mgals (Biehler, in

Fig. 5. Faults and geothermal areas in the Imperial Valley. Contours indicate near-surface temperature gradients ($^{\circ}\text{C}/100\text{ m}$) measured in more than 100 shallow boreholes, using data up to 1972. Dot pattern indicates basement rock outcrops. (Source: Elders, *et al.*, 1972.)

Rex, *et al.*, 1971). The excess mass in the subsurface has been attributed, at least in part, to thermal metamorphism of the sedimentary rocks by the hot brines. For example, intense metamorphism of the sedimentary fill occurs in the Salton Sea geothermal field. Active formation of greenschist facies rocks is occurring at depths of 1 to 2.5 km below the surface, where the temperature ranges from 300 to 350°C (Muffler and White, 1969). Brines recovered from these depths contain up to 25 wt% of total dissolved solids (Helgeson, 1968).

Such high temperatures and highly saline brines have not been found in the other thermal anomalies drilled to date. Temperatures from 150 to 250°C and brines containing from 3,000 to 20,000 ppm of total dissolved solids are much more characteristic. Examples from Cerro Prieto, East Mesa, and Heber are discussed later in this report. Similarly, the degree of metamorphism observed is characteristically less than that seen in rocks from the Salton Sea field.

In general, there is no surface expression of the thermal anomalies, except in the Cerro Prieto and Salton Sea fields. In these fields there are, or were, associated warm springs, mudpots, and Quaternary volcanoes (Fig. 3). The only other anomaly known to be directly associated with igneous rocks is at Heber, where one drill hole penetrated an approximately 15-m-thick basalt sill (D. Butler, personal communication, 1974). The much larger-than-average gravity highs at the Cerro Prieto and Salton Sea fields are presumed to be associated with igneous intrusions at depth.

Surface expression of the thermal anomalies is retarded by impermeable caprocks. For example, the Salton Sea geothermal field has an impermeable caprock of lacustrine clays up to 450 m thick (Helgeson, 1968; Randall, 1974). The Dunes hydrothermal system, on the other hand, has an impermeable caprock developed by self-sealing. In the upper 300 m of a 612-m deep borehole in this anomaly, there are seven intervals of intense sandstone-to-quartzite cementation, with densities as high as 2.55 g/cm³ and porosities as low as 3% (Elders and Bird, 1974; Bird, 1975).

A recently completed study of the Salton Sea anomaly, based on subsurface samples and well logs, shows the structure of the field to be complex (Randall, 1974). While it is possible to correlate stratigraphic horizons in a north-south direction, structural discontinuities (presumably faults) make correlation impossible in an east-west direction. Perhaps the most surprising finding of this study is that there is no obvious relationship between major stratigraphic and lithological horizons and the temperature field. Similarly, the isothermal surfaces appear to ignore major structures. Randall concludes, therefore, that the heat source is a young intrusion and that the distribution of temperatures is, at present, controlled almost entirely by distance from this cooling intrusive body.

VOLCANISM

Volcanoes at the Cerro Prieto and Salton Sea fields are apparently part of the suite of volcanic activity associated with the East Pacific Rise and the Gulf of

California. The Barcena volcano in the Revillagigedo Islands on the East Pacific Rise (R in Fig. 4) erupted in 1952. Basalt dredged from the deep basins in the Gulf is olivine tholeiite similar to that found on the East Pacific Rise and other ocean spreading centers (Hawkins, in Elders and Biehler, 1975). No detailed studies of the Cerro Prieto volcano have been published; however, it is a lithoidal rhyodacite cone, which appears to be the product of a single eruptive cycle. The marked lack of erosion of the cone attests to the relatively youthful age of this eruption. It is on this basis that I have indicated it in Fig. 3 as being Quaternary in age.

At the south end of the Salton Sea are five small extrusive rhyolite domes arranged along a northeast trend. These domes, collectively known as the Salton Buttes, were extruded onto Quaternary alluvium. A single K-Ar age determination on the westernmost dome, Obsidian Butte, gave an age of approximately 16,000 years (Muffler and White, 1969). Two of the domes, those at Red Hill, are linked by subaqueous pyroclastic deposits; the others are single extrusions with or without marginal lava flows. All of the domes consist of low-calcium, alkali rhyolite with only 1 or 2% crystals. Similar rocks recovered from geothermal wells had been altered extensively by water-rock reactions. The fresh rhyolites are identical in composition to soda rhyolites erupted on the islands of the East Pacific Rise (Robinson, Elders, and Muffler, 1975). Basaltic rocks occur as xenoliths in the domes and as subsurface dikes, sills, or flows. Except where hydrothermally altered by brines, these

rocks are also identical to low-potassium tholeiitic basalts erupted on the East Pacific Rise and on islands in the Gulf of California (Robinson, Elders and Muffler, 1975). These observations support the hypothesis that the conditions that control magma genesis under the Salton Trough and the Gulf of California are similar to those operating beneath oceanic spreading centers.

Numerous partly melted granitic xenoliths in these rhyolite domes show various degrees of either cotectic melting along quartz-feldspar boundaries or disequilibrium incongruent melting of hydrous-ferromagnesian minerals. These granite inclusions contain notably higher SiO_2 , CaO , and Na_2O and lower total iron than the enclosing rhyolite. The compositions and textures of these rocks suggest that they are fragments of the basement rather than being cogenetic with the rhyolites. This bimodal basalt-rhyolite assemblage in the Salton Sea geothermal field is believed to have formed in two stages by partial fusion of mantle peridotite, forming successive rhyolitic and basaltic melts. After formation, the rhyolite magma was partly contaminated by continental crust material.

A compound, positive, magnetic anomaly only in part due to the exposed rhyolites is associated with the Salton Sea geothermal field. A long magnetic high, 5 to 8 km wide, is centered on the southern half of the lake and extends 28 km in a northwesterly direction. Griscom and Muffler (1971) interpret this high as being due to intrusive rocks at depths greater than 2 to 2.5 km. These rocks may take the form of a stock-sized pluton or a concentrated dike swarm. Also,

associated with the magnetic and heat-flow anomaly is a residual-gravity anomaly of approximately 22 mgals. The amplitude of the gravity anomaly suggests a center of gravity for the excess mass at about 6 km depth (Biehler, personal communication, 1974). Therefore, there are several lines of evidence indicating that the heat source for this anomaly is an igneous intrusion. A combined geological, geophysical, and geochemical investigation is underway at the University of California at Riverside to investigate the Salton Sea geothermal field much more thoroughly.

TECTONICS

The crust in the Salton Trough is being actively deformed. Since triangulations began in 1931, there appears to have been about 2 m of differential, right-lateral movement in the southern part of the Imperial Valley. Similarly, there have been downward level changes of tens of centimeters during this period. The most notable deformations of the survey net are associated with the May 1940 earthquake on the Imperial Fault (7.1 magnitude on Richter scale). Clearly, the Imperial Valley as a whole is undergoing right-lateral horizontal motion and the center is subsiding relative to the walls by steady creep punctuated by earthquake activity (Elders, *et al.*, 1972).

Viewed in context of regional tectonics, these modern earth strains provide a good model for the origin and history of the Trough. Overall, the region is characterized by a tectonic environment dominated by strike-slip, or transform, faulting. However, at points where the

transform faults are en echelon, tension gaps and compression zones are formed. The tension gaps, termed "rhombochasms" by Carey (1958) or "pull-apart basins" by Crowell (1974) predominate. These are sites of high heat flow and low topography where localized crustal spreading can occur. It is in these areas that new material can be introduced by leakage of magmas from the mantle as the crust thins (Elders, *et al.*, 1972). Garfunkel (in Rex, *et al.*, 1972) suggests that depressed areas on either side of the apex of the Colorado River delta may be subtle expressions of such tectonic subsidence that have not been obliterated by sedimentation. The en echelon arrangement of the western boundary of the Imperial Valley, with its numerous low embayments, also supports this idea. Figure 6 shows a possible arrangement of tension "pull-apart" basins, compression zones, and strike-slip faults. For example, young volcanic rocks crop out at Consag Rock in the Wagner Basin (W in Fig. 4), the northernmost of the closed basins in the gulf of California. Thatcher and Brune (1971) have described an earthquake swarm that occurred in this basin in 1969. Over 70 shocks, with magnitudes of between 4.0 and 5.5 (Richter scale), occurred during a 6-hr period. At the same time, there appeared to be a sympathetic coupling with earthquakes in the next basin to the south, the Delfin Basin (D in Fig. 4). These two basins appear to be related by a transform fault system. Lomnitz, *et al.*, (1970) and Elders, *et al.*, (1972) have suggested that this pattern of transform faults and "pull-apart" basins persists to the north. The location of Cerro Prieto between the San Jacinto

Fig. 6. Possible relationship between pull-apart basins and strike-slip faulting in the Salton Trough. Postulated "spreading centers" or tensional zones, young volcanics, geothermal areas, and zones of intense folding and compression in Tertiary sediments are indicated. (Source: Elders, et al., 1972.)

and the Imperial faults fits this model. Similarly, the Brawley and Salton Sea fields occur between en echelon fault systems. A similar relationship between major through-going faults and the other geothermal fields has not been established. However, the Heber and East Mesa anomalies are elongated perpendicular to local faults. Throughout much of the Imperial Valley the rapid rate of sedimentation obscures faulting, and few surface traces of faults can be positively identified. However, in addition to the major faults, which are visible at outcrop, others have been inferred from infrared aerial photographs (Babcock, 1971), from seismic reflections (Sigurdson, et al., 1971), from electrical resistivity (Meidav and Ferguson, 1972), and from micro-earthquake studies (Combs and Hadley, 1974). When the structure of the Valley is more thoroughly understood, a clearer relationship between geothermal areas and faults may emerge.

GEOHERMAL RESOURCES

A great deal of exploration and assessment work is being carried out north of the U.S.-Mexico border, whereas the only significant electrical power production is at Cerro Prieto. Estimates of total Salton Trough geothermal resources cannot be made with any degree of certainty at this time, and published estimates vary widely. For example, the U.S. Geological Survey has estimated total

usable and recoverable water in the Imperial Valley to be about 250 km³ of water at 150°C or hotter (Dutcher, et al., 1972). In defining usable water, these authors included only water containing less than 35,000 mg/liter of dissolved solids. Thus, the hypersaline brines of the Salton Sea field were excluded from this estimate. An earlier estimate of water resources, using the same definition, suggested that as much as 6,000 km³ of water might be recoverable (Rex, 1970). Moreover, efforts are underway to develop the technology to handle hypersaline brines. Obviously, better information is needed to improve resource estimates.

The geothermal resources of specific thermal anomalies are discussed later in this report. Geothermal investigations in the Imperial Valley are being conducted by a diverse group, including utility companies, major oil companies, small private companies, universities, and government agencies at the federal, state, and county level. There is no centralized authority to coordinate these efforts. The rapid pace of development coupled with the release of proprietary information already developed by private industry could change our understanding of the origin and nature of these thermal anomalies at any time. Cooperation between the various disparate groups involved already is occurring to the mutual advantage of everyone concerned.

History: Salton Trough Geothermal Development*

The existence of geothermal activity in the Salton Trough has been known for many years. Live steam fumaroles, mud volcanoes, and boiling mudpots (Figs. 7 and 8) were observed by the area's earliest inhabitants in the Salton Sea near the volcanic outcrop now known as Mullet Island. The first attempts at resource recovery began here and, in following years, spread to other areas of the Salton Trough: Brawley, East Mesa, Heber, and Cerro Prieto.

The first exploratory geothermal wells were drilled on Mullet Island in 1927 by the Pioneer Development Company (Fig. 9). Three wells were drilled into this volcanic dome, the deepest to a depth of 449 m. Steam and hot water were found in all three wells, but pressure and water volume were not sufficient for commercial operation so the wells were abandoned. However, in drilling these wells, it was noted that a large volume of carbon dioxide gas was produced. This observation led to the discovery of the adjacent Imperial Carbon Dioxide Field, which, from 1933 to 1954, supported some 55 producing wells. The carbon dioxide was recovered from shallow sands (150 to 200 m deep) containing hot water (as high as 66°C). In conjunction with the field, two processing plants were built in the area to convert the carbon dioxide to dry ice.

The first exploratory well to produce substantial amounts of steam and hot water was the Kent Imperial Corporation's

Sinclair #1, located near Niland in the Imperial Valley (Fig. 10). Originally an oil prospect, this well was drilled in 1957-58 to a depth of 1440 m. However, when tests of the most promising oil zone produced only hot water and steam, the site was prepared for steam production. In 1959, a small pilot plant (including separators, condensers, and generators) was installed at the well. The plant operated intermittently for 4 months before being shut down when the well scaled up at the surface. The well has not produced since.

From 1961 to 1964, a number of geothermal wells were drilled in the Salton Sea area. The first was the Joseph I. O'Neill, Jr., Sportsman #1 about 6 km northeast of Sinclair #1 (Fig. 9). This 1441-m-deep well proved to be a good steam producer. As a result, 10 additional wells were drilled in the immediate vicinity, 8 of which also became successful steam producers. All of these wells produced steam from hot water reservoirs flowing a mixture of steam and high-salinity water at the wellhead. Brine analyses showed surprisingly high mineral contents, some of the brines having total dissolved-solids concentrations of more than 300,000 ppm. This high salinity, accompanied by some caustic properties, caused severe corrosion and scaling problems in the wells and support equipment. Recognizing the potential for mineral recovery from the brines, Imperial Thermal Products, Inc. (a subsidiary of

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Fig. 7. Mud pots with carbon dioxide vents near Mullet Island, Salton Sea area, California.

Fig. 8. Mud volcanoes near Mullet Island, Salton Sea area, California.

Morton International, Inc.) and the Earth Energy Co. (a subsidiary of Union-Pure Oil Co.) both located pilot power-production/mineral-recovery plants in the area. However, after several years of experimentation, these ventures were terminated as being uneconomical.

Also during this time period (1963), the Standard Oil Company of California drilled a 4,097-m-deep oil-exploration well in the vicinity of Brawley. When abnormally high temperatures were encountered, it prompted the Oil company to drill 17 shallow, temperature-test

Fig. 9. Salton Sea Geothermal Field, geothermal well locations.

Fig. 10. Kent Imperial Corporation geothermal well Sinclair #1.

holes in the southern Imperial Valley. When these test holes also indicated temperature anomalies, it greatly increased interest in the area's geothermal potential.

In 1968, primarily through the efforts of R. W. Rex (a former Standard Oil Company employee), the University of California at Riverside (U.C. Riverside) began an intensive geothermal investigation of the Imperial Valley. This program, supported by the U.S. Bureau of Reclamation, the National Science Foundation, Standard Oil Company of California, the Chevron Oil Field Research Company, and the Imperial Irrigation District, incorporated Standard Oil's previous work and included heat-flow measurements; resistivity, gravity, and seismic surveys; and more than 100 shallow-hole temperature tests. The study located several significant geothermal anomalies (Fig. 11), most of which have since been classified

as Known Geothermal Resource Areas (KGRA), a Federal designation based on temperature gradients (Rex, *et al.*, 1971), and other geophysical data.

During 1972, 12 wells were drilled in the Salton Trough: 5 in the Salton Sea area, 3 on the Heber anomaly, 1 on the East Mesa anomaly, and 1 on the Dunes anomaly. All but two of these wells were completed as potential steam and hot-water producers. Also in 1972, steps were taken to establish a subsidence surveillance network in the Imperial Valley and adjacent lands; surface subsidence was considered a distinct possibility in light of the large-scale withdrawal of geothermal fluids. The network, coordinated by the California Division of Oil and Gas and funded primarily by the National Geodetic Survey, the National Science Foundation, the U.S. Bureau of Reclamation, Imperial Irrigation District, the County of Imperial, and the U.S. Dept. of Transportation, was first surveyed in 1971 and 1972. Natural surface movements were determined to provide a basis by which to detect movements due to fluid withdrawal. The network was resurveyed in 1973 and 1974, and is to be resurveyed every two years—more frequently, if extensive geothermal production is anticipated.

During 1973 and 1974, a total of 12 geothermal wells were drilled in the Imperial Valley: 4 in the central valley, 4 at East Mesa, 1 in the Salton Sea area, and 3 in the Heber area. In addition, several ongoing experimental programs were initiated to test and evaluate methods, materials, and equipment for the handling and use of geothermal fluids. Three of these projects are being conducted in the Salton Sea area, one by the San Diego

Fig. 11. Temperature gradient map of the Imperial Valley, California, showing KGRAs.

Fig. 12. Union Oil Company geothermal well Veysey #1, 4 km north of Brawley, California.

Gas and Electric (SDG&E) Company, another by the Phillips Petroleum Company, and a third by the Lawrence Livermore Laboratory. A test program is also underway in the Heber area by the Chevron Oil Company. Finally, the U.S. Bureau of Reclamation is investigating the use of

geothermal brines in two experimental desalination plants at East Mesa.

In the Mexicali Valley, the Mexican Government is operating a 75-MW geothermal power plant. Electricity generation began in April 1973, culminating 13 years of development starting with exploratory drilling in 1961 and plant construction in 1968.

In early 1975, the Union Oil Company began drilling Veysey #1, a geothermal well on the Brawley anomaly about 4 km north of the town of Brawley (Fig. 12). Four additional wells are proposed for the same general area. These sites are near the center of the geothermal-gradient anomaly reported by the U.C. Riverside and the U.S. Bureau of Reclamation in their joint study published in 1972. These Union Oil Company wells will be the first field tests of the Brawley KGRA.

Finally, as mentioned earlier, the Morton Salt and Union Oil Companies have tried unsuccessfully to develop economic processes for large-scale mineral extraction from geothermal brines. Nonetheless, mineral-recovery projects continue. For the past several years, the Geothermal Energy and Mineral Corporation has been extracting calcium chloride from the brine of Sinclair well #4. And a similar, small-scale calcium chloride operation is underway at the Morton Salt Company's IID #1 well.

More-detailed descriptions of the geothermal research and development projects mentioned here are presented in the "Industrial and Institutional Geothermal Projects" section of this report.

Industrial and Institutional Geothermal Projects

IMPERIAL VALLEY*

The U.C. Riverside has been instrumental in developing the geothermal resources of the Imperial Valley. In 1968, a group of faculty and students, headed by R.W. Rex, began a series of geothermal study projects at the area that continue today. This program, supported by the National Science Foundation, the U.S. Bureau of Reclamation, the Academic Senate of U.C. Riverside, the Standard Oil Company of California, the Chevron Oil Field Research Company, and the Imperial Irrigation District, includes geological and geochemical studies, geothermal-gradient and heat-flow measurements, and resistivity, gravity, and marine seismic surveys. Since its inception, this field research program has identified five geothermal anomalies and a fault system in the area. Moreover, four of these anomalies were included in the U.S. Geological Survey's 1971 list of KGRA's (Fig. 8). During 1972, the California Department of Water Resources, in a joint effort with U.C. Riverside and the U.S. Bureau of Reclamation, drilled a 612-m well on the Dunes anomaly. The Bureau of Reclamation also began a drilling program on the larger East Mesa anomaly the same year.

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EAST MESA†

At East Mesa (Fig. 13), the Bureau of Reclamation is exploring the feasibility of desalting geothermal brines and the practicability of concurrently generating electric power. Additional fresh water developed in the Imperial Valley can be used to augment flow in the Colorado River system, the most highly regulated and intensively utilized river system in the U.S. The recovery of marketable mineral products from blowdown brines is another possibility being investigated in this multipurpose research and development program.

A total of 60 test holes have been drilled into the East Mesa anomaly, and detailed temperature profiles have been measured in each. Based on these data, heat-flow values and contours have been computed, outlining the lateral extent of the Mesa anomaly (Fig. 14). Gravity, microearthquake, seismic noise, and resistivity methods have been used to further define the anomaly. To date, five test wells, each more than 1829 m deep, have been completed at the site. Bottom temperatures at these wells range from 150 to 200°C. Mesa 6-1 and Mesa 6-2 are production wells. Mesa 5-1 is being developed as an injection well, based on surface and subsurface geophysical data.

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Fig. 13. Bureau of Reclamation East Mesa test site, Imperial Valley, California. Geothermal test facility and well Mesa 6-1 in foreground; brine holding pond in background. (Photo courtesy of U.S. Bureau of Reclamation.)

The other two wells, Mesa 31-1 and Mesa 8-1, are to be tested as production wells.

Two experimental desalination plants, a multistage flash-distillation unit (MSF) and a vertical-tube evaporator-distillation unit (VTE), are being used by the Bureau of Reclamation to test and evaluate procedures for desalting geothermal fluids. Each unit is designed to produce 75 to 190 kl of distilled water per day, depending on operating conditions.

Distillation-type desalination methods are particularly attractive for geothermal brines because the fluids are naturally

heated. To date, both plants have operated successfully and have produced tens of kiloliters of distilled water. The VTE has been run continuously in around-the-clock operations, with no evidence of scaling or corrosion problems in the heat exchanger tubes. Operational data from these test desalination units will provide design criteria for larger, prototype units as well as information for the design of heat exchangers to be used in geothermal power-plant.

The East Mesa facility provides a setting where geothermal research and

Fig. 14. Heat flow, U.S. Bureau of Reclamation Test Site, East Mesa, Imperial Valley, California. (Figure courtesy of Bureau of Reclamation.)

development can be performed under actual field conditions using producing geothermal wells. Therefore, the facility has been made available to scientists studying the characteristics of geothermal fluids.

MEXICALI VALLEY*

The southern extension of the Imperial Valley, the Mexicali Valley, is under geothermal development by the Mexican Government (Comision Federal De Electricidad). About 2.5 km² of the Cerro

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Fig. 15. Geothermal well and water-steam separator, Cerro Prieto power plant, Mexicali Valley, Baja, California.

Prieto geothermal field (about 32 km south of Mexicali, Baja, California) currently is being exploited by 30 steam/water-producing wells (Fig. 15). The reservoir consists of alternating shale and sandstone layers with a cap rock of impermeable clay. The wells, about 1300 m deep, use 7-5/8-in. production casing. Maximum bottom-hole temperature is about 350°C, with a maximum shut-in pressure of about 9 MPa. Average steam production is about 55,000 kg/hr.

After an extensive field-test program, the Mexican Government in 1968 started construction of a 75-MW geothermal power plant in the Mexicali Valley (Fig. 16). Commercial operation began in April 1973. The plant uses two Toshiba turbines (Figs. 17 and 18) in two 37,500-kW turbo-alternator groups, having an admission pressure of 517 kPa and an exhaust-steam pressure of 12 kPa. The facility operates on steam from 13 geothermal wells. Steam condensers are of the barometric type (Fig. 19), with ejectors to extract incondensable gases (mainly carbon dioxide and hydrogen sulphide). As this is a hot-water reservoir, the separated saline water must be disposed of in evaporation ponds (Fig. 20). Salts from this discharge water are being considered as potential sources of marketable chemical products (e.g., potassium chloride and lithium carbonate).

The total investment for the first stage of this project was \$20 million, nearly \$300 per kilowatt of installed capacity. The total cost per generated kilowatt-hour is about 8 mills, including operation cost, amortization, depreciation, maintenance cost, and the cost of new wells to sustain present power generation.

Fig. 16. Cerro Prieto geothermal power plant, Mexicali Valley, Baja, California.

Fig. 17. Toshiba 37,500-kW turbo-alternator groups, Cerro Prieto power plant, Mexicali Valley, Baja, California, Mexico.

HEBER*

Geothermal development in the Heber area has been due largely to the efforts of the Chevron Oil Company, Magma Power, and associated organizations. This area has been considered a potential heat anomaly since 1945, when Amerada drilled an oil and gas test well (Timkin #1) just west of the Heber townsite. Well logs indicated a higher-than-normal heat gradient. In 1963, Chevron confirmed the Heber geothermal anomaly by drilling a shallow test hole and measuring the heat gradient. In 1971, Chevron initiated a geophysical program to more fully evaluate the geothermal potential of the Heber area.

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Fig. 18. Toshiba turbine rotor, Cerro Prieto power plant, Mexicali Valley, Baja, California.

Fig. 19. Barometric steam condensers, Cerro Prieto power plant, Mexicali Valley, Baja California.

Early in 1972, Magma Energy, Inc., with the financial support of Chevron Oil Company, drilled the Hultz well #1 to a depth of 1,569 m. This and subsequent

Fig. 20. Waste-water discharge to solar evaporation ponds, Cerro Prieto power plant, Mexicali Valley, Baja, California.

wells, are shown in Fig. 21. Well conditions were encouraging and prompted Magma to drill a second, 1,524-m-deep well nearby. In the fall of 1972, Chevron drilled its first well, Nowlin Partnership #1, to a depth of 1,533 m.

In 1973, Chevron, New Albion Resources Company (a subsidiary of San Diego Gas and Electric Company) and Magma Energy agreed to a joint venture at Heber to test and evaluate the geothermal potential indicated in the earlier wells. Chevron was designated operator for the joint venture. The resulting test project began in late 1973 and was completed in mid-December 1974. The objective of this program was to determine production and injection capability of the Heber reservoir and to evaluate binary-cycle heat-exchange technology. Chevron's Nowlin Partnership #1

Fig. 21. Geothermal wells in the Heber area, Imperial Valley, California.

and Magma's Hultz well #1 were produced, and reservoir fluids were injected into Magma's Hultz well #2.

During the summer and fall of 1974, Chevron drilled two additional wells for the joint venture and one of its own. Test data from these wells will be evaluated with that from other wells to determine if reservoir conditions warrant full-scale commercial development.

Because the Heber geothermal anomaly is in the moderate temperature range, the fluid must be pumped to enable a sufficiently high temperature to be maintained at the surface. The method anticipated to be applicable at Heber is a binary process using heat exchangers (Fig. 22).

Recently, San Diego Gas and Electric has installed a heat-exchanger test module at the Chevron site (Fig. 23) and is collecting information on heat-exchanger

Fig. 22. Schematic of binary process for geothermal power generation.

performance. Heber fluids have much lower salinity than Salton Sea fluids (about 10% of the total dissolved solids); therefore, direct heat-exchanger operations may be feasible with Heber fluids.

Fig. 23. Chevron Oil Company geothermal test site, Heber area, Imperial Valley, California.

SALTON SEA GEOTHERMAL FIELD*

Wells in the Salton Sea geothermal area usually encounter hot water (150 to 300°C) between 300 and 1500 m below the surface. The concentration of solids in this water generally is 20 to 30 wt%. About 20% of the water flashes into a steam-water mixture as it flows up the well pipe, leaving a highly concentrated salt brine. These conditions impose severe drilling and production problems on any attempt to exploit the resource. Although large volumes of heated, highly mineralized water are now known to exist in this area, commercial exploitation has been largely unsuccessful primarily due to the highly corrosive nature of the brines plus the problems of residual-salt disposal. In

other respects, however, reservoir characteristics are attractive enough to encourage ventures to overcome these obstacles. The reservoir occupies an estimated 65 km², and sustained production yields for the few wells tested have reached 227,000 kg/hr of brine. Such energy potential, located close to large population centers, has maintained interest in this geothermal area. Some of the more ambitious attempts to develop Salton Sea resources are summarized in the following subsections.

Imperial Thermal Products—Morton IID Wells*

Early in 1965, Imperial Thermal Products (ITP), a subsidiary of Morton International, Inc., contracted with various

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geothermal well owners in the Salton Sea area to investigate the technical problems associated with the commercial use of geothermal brines. The ensuing ITP research and development program focused on two areas:

- Process steam for power generation
- Process brine for mineral extraction.

A 3,000-kW steam-operated power plant was constructed at IID well #1, and solar

evaporation ponds were located nearby (Fig. 24). Token amounts of electrical power were generated and delivered to the local Imperial Irrigation District. However, brine-salt carryover in the steam phase caused severe scaling and corrosion of the turbine blades and auxiliary equipment. Power generation was eventually discontinued for further research into the scaling problem.

Two different techniques were considered for mineral recovery. One, steam

Fig. 24. Imperial Thermal Products, Inc., Morton IID Well #1, Imperial Valley, California. (Photo courtesy of Morton International, Inc.)

and super-heated brine could be processed through vacuum pans to precipitate out recoverable minerals, with the steam then being used for power generation. Or, two steam could be flashed from the brine and the brine then discharged into open ponds, where marketable minerals could be concentrated by solar evaporation and fluid leaching. The second method was selected for full-scale testing because of its lower production cost and greater steam yield. Much of the original research and development of mineral extraction and harvesting methods is accredited to ITP. Calcium chloride, potassium chloride, and other mineral salts were successfully extracted and marketed by the Company until declining prices of competitive sources eventually brought an end to the extraction venture.

Other ITP research efforts included: reservoir characterization, brine production and control, brine content analysis, and reinjection techniques. The ITP site is no longer in operation, but it still represents the first successful attempt at generating electrical power using Salton Sea geothermal brines.

Earth Energy Company—River Ranch Well #1*

In 1964-65, Earth Energy Company (a subsidiary of Union-Pure Oil Co.) located a geothermal pilot plant at the River Ranch well #1 in the Niland area (Fig. 25). Solar evaporation ponds were constructed on adjoining property, and a pilot separation plant was erected at the well. Attempts to produce electrical

power and extract minerals from the brine met with limited success. Operation of the wells ultimately shifted to Union Oil, who expanded research operations temporarily and then discontinued the Niland operation in favor of the Brawley geothermal area. Union presently plans to drill four exploratory wells in the Brawley area.

In 1963, the Colorado River Basin Regional Water Quality Control Board prohibited the discharge of any geothermal brines into any channel draining into the Salton Sea. This prompted Earth Energy and Imperial Thermal Products to experiment with brine injection wells as an alternative to surface disposal. In 1965, surplus brine from River Ranch well #1 was injected into Union Oil's Hudson Ranch well #1 and brine from IID well #1 was injected into Morton's IID well #3. Reinjection is now generally considered to be the best method of waste-brine disposal and continues to be an area of active research.

Fig. 25. Earth Energy Company's geothermal well River Ranch #1, Imperial Valley, California.

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Phillips Petroleum Company – Sinclair Wells *

The Sinclair wells (located on the Sinclair Ranch) are among the earliest efforts to exploit the Salton Sea's geothermal resources. Sinclair well #4 and 77 acres of solar evaporation ponds (Figs. 26 through 28) were used as early as 1966 by the Chloride Products Company, Inc., and Western Geothermal, Inc., to produce calcium chloride. Operations were later conducted by Hills Brothers, and presently are in charge of Lee Chemical.

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The current Sinclair-wells geothermal operation was organized in 1972, as a joint venture of the Southern Pacific Land Company, Southern California Edison Company, and Phillips Petroleum Company. The basic plan of this joint research and development program is to:

- Establish a reliable, continuous, geothermal brine production and injection disposal system
- Investigate scaling and corrosion problems and develop methods to minimize them
- Perform studies aimed at utilizing geothermal brines for power production.

Fig. 26. Phillips Petroleum Company geothermal well Sinclair #4, Imperial Valley, California.

Fig. 27. Wellhead, geothermal well Sinclair #4, Imperial Valley, California.

With these goals in mind, two of the Sinclair wells were reconditioned and an

Fig. 28. Test site, geothermal well Sinclair #4, Imperial Valley, California.

inter-well fluid line was designed, installed, and tested. Three periods of well flow have been completed to establish the nature of the fluids, determine well characteristics, identify hydraulic-mechanical relationships, and conduct dynamic system tests. This operation is still underway, and an expanded program is planned.

San Diego Gas and Electric
Company—Magma
Power Company Wells*

San Diego Gas and Electric (SDG&E) entered the geothermal arena in 1971.

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Two factors contributed heavily to this action. First, was the realization that the availability of low-cost natural gas to fuel power plant boilers was diminishing rapidly. And second, was the fact that work underway in the Imperial Valley by U.C. Riverside, under the sponsorship of the U.S. Bureau of Reclamation, was demonstrating a potential geothermal resource within an economic distance of the Company's service area (San Diego County and a portion of southern Orange County).

The Utility actually became involved in geothermal activities through an agreement with its wholly-owned resource subsidiary, New Albion Resources Company (NARCO) and the Magma Power Company and Magma Energy, Inc. The agreement provided that, in return for funding drilling and testing on Magma leases in the Imperial Valley, NARCO could obtain an interest in the leases.

During early 1972, exploratory drilling was conducted in several areas of the Valley to identify geothermal sites for development. The most promising locations were the Niland area, at the south end of the Salton Sea, and the Heber area, south of the town of El Centro. A second well was drilled in each of these areas to confirm initial findings and to provide reinjection capability for flow testing.

Flow testing was carried out at the Niland site in mid 1972, using a large wellhead separator on loan from the Union Oil Company. The test provided information on well deliverability, fluid characteristics, and other information necessary to design a geothermal test facility to evaluate the feasibility of power generation from the resource.

In August 1972, the C. F. Braun Company was contracted to design and procure major equipment for a geothermal test facility. The facility was designed on the binary system, using a single steam flash, with the steam and brine from the flash tank being directed through heat exchangers that, in turn, heated isobutane. This facility was completed in early 1973. An area of uncertainty in the design was the longevity of the brine exchangers. Therefore, the Braun Company decided to run some small-scale field tests prior to starting construction of the geothermal test facility.

During 1973, a small-scale test facility was constructed and installed at the Niland site. A flow diagram of the 1973 test hardware is shown in Fig. 29. Fluids from the producing well entered the separator at 150 psig. Steam from the top of the separator passed through one heat exchanger, and brine leaving the bottom of the separator flowed through another heat exchanger. Both steam and brine temperatures were about 190°C. Mineral content of the geothermal brine was approximately 200,000 ppm. Concentrations of solids in the steam leaving the test separator ranged from 40,000 to 80,000 ppm. In both the steam and brine exchangers, heat transfer performance declined from an initial value to design limits after some 100 hr of operation (due to scaling). Testing continued until the fall of 1973; however, a satisfactory solution was not found for mitigating the brine and steam scaling problem. Therefore, construction of the large geothermal test facility was delayed.

In early 1974, work resumed to develop effective methods for separating steam from geothermal brine and scrubbing the

Fig. 29. Schematic of 1973 geothermal field test system: heat exchanger system, San Diego Gas and Electric Company. (Figure courtesy of SDG&E.)

steam to achieve as pure a product as possible. Brine heat-exchanger testing was discontinued. Figure 30 shows a flow diagram of the 1974 test apparatus. These tests, completed during the summer of 1974, consisted of passing steam through separators, scrubbers, and heat exchangers to determine if scrubbing the steam would eliminate the severe scaling problems.

The first phase of the 1974 test program consisted of evaluating different types of steam separators. The separator selected yields outlet steam with a solids content of approximately 200 ppm, a considerable reduction from the initial well fluid solids content of 200,000 ppm. Figure 31 is a cutaway view of the separator, showing the interior of the vessel. Both the first-stage (150 psig) and second-stage (50 psig) separators have the same configuration.

Well fluid enters the separator through the port in the bottom left side of the chamber and impinges on the vessel end dome, where a plate protects the vessel wall. Steam leaves the separator through a port in the upper right corner of the vessel. Brine collects at the bottom of the separator then flows to the second-stage unit through the port in the lower right corner of the vessel.

The interior of the steam scrubber is shown in Fig. 32. Steam from the separator enters the scrubber through the lower left-hand port. It then flows upward through five trays of pure water, which scrubs solids out of the steam. The cleaned steam then exists at the top of the scrubber. Wash water is added continuously to the scrubber through the top of the vessel, cascading down to exit drains

Fig. 30. Schematic of 1974 geothermal field test system: two-stage flash system, San Diego Gas and Electric Company. (Figure courtesy of SDG&E.)

at the bottom. During the 1974 test, solids content in the scrubbed steam ranged from 10 to 20 ppm, compared to the 100 to 200 ppm entering the scrubber from the separator.

Fig. 31. C.F. Braun scale-model separator, 1974 field test, San Diego Gas and Electric Company. (Figure courtesy of SDG&E.)

Fig. 32. Ben Holt steam scrubber, 1974 field test, San Diego Gas and Electric Company. (Figure courtesy of SDG&E.)

The type of heat exchanger used in the 1974 tests is shown in Fig. 33. The first-stage heat exchanger employed steam flowing on the outside of the tubes while the second-stage unit used steam flowing through the tubes. Approximately 454 kg/hr of steam (for both first and second stages) flowed through the heat exchangers and was condensed. A closed loop of distilled water circulated through the heat exchanger provided the cooling mechanism.

The overall heat transfer coefficient versus operating time for the 1973 and 1974 first-stage tests is shown in Fig. 34. The 1974 test results indicate that the heat exchangers will operate for 3,200 hr before reaching design conditions requiring cleanup, an encouraging improvement over the 108 hr required to reach cleanup conditions in the 1973 tests. Total operating time for the first-stage heat exchanger during the 1974 test was 398 hr, sufficient time to establish a trend on which to predict the number of operating hours before cleanup would be required.

Figure 35 shows similar data for the second-stage steam heat exchangers. Here, 1973 test results indicated 81 hr to reach design conditions requiring cleanup. On the other hand, 1974 test data indicates that the second-stage heat exchangers will operate to 10,050 hr before needing cleaning. Operating time for the second-stage heat exchangers during the 1974 test was 587 hr.

A flow diagram of the anticipated, full-scale, geothermal test facility is shown in Fig. 36. Unlike the two-stage tests, this facility will not use the brine in the heat exchangers, but will flash it into steam in three stages. Then, after passing through scrubbers, the steam will be routed through heat exchangers to heat an isobutane working fluid that, in turn, will be used to power a turbine generator. After use, the steam will be recondensed, combined with residual brine, and discharged through reinjection wells. The Ben Holt Company has been contracted by SDG&E to do the necessary redesign work

Fig. 33. Heat exchanger, 1974 field test, San Diego Gas and Electric Company. (Figure courtesy of SDG&E.)

Fig. 34. First-stage heat exchanger tests, San Diego Gas and Electric Company. (Figure courtesy of SDG&E.)

Fig. 35. Second-stage heat exchanger tests, San Diego Gas and Electric Company. (Figure courtesy of SDG&E.)

Fig. 36. Three-stage steam flash system, San Diego Gas and Electric Company. (Figure courtesy of SDG&E.)

on the geothermal test facility. Construction of the facility is expected to start in 1975.

LLL Geothermal Program — Salton Sea Geothermal Field*

In January 1974, the Lawrence Livermore Laboratory has initiated a geothermal research program aimed at developing methods for the recovery and conversion of geothermal energy in the high-temperature, high-salinity (up to 30% total dissolved solids) brines known to exist in the Salton Sea Geothermal Field. Toward this end, a small experimental

power plant (<10 MWe) is planned to evaluate the technical and economic feasibility of new concepts and to provide a continuing source of operating data to aid in the design of commercial-size systems. Successful exploitation of Salton Sea geothermal resources depends largely on solving the myriad of technical problems associated with handling brine, controlling scale and precipitation, and designing corrosion/erosion-resistant systems for efficiently converting geothermal energy to electrical power.

One method used to convert geothermal energy to electricity is the flash-steam system (Fig. 37). In this process, a well-head mixture of hot water and steam is piped into separators where the steam

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Fig. 37. Methods for electric power generation from hot-water deposits.

fraction is separated from the liquid. The saturated steam, then, is used to drive standard multistage turbines, and the remaining hot liquid is discarded. However overall thermal efficiency of this process, using brines at 260 to 315°C, is only about 11%. Moreover, the flash-steam method generally is limited to low-salinity water (< 2%). With higher-salinity brines, salt carryover in the steam causes scaling in the separator.

In an attempt to solve these difficulties and to provide a more thermally efficient conversion process, a binary-cycle system (Fig. 37) is being developed in the geothermal industry. This system uses a heat exchanger to transfer the thermal energy to a clean, low-boiling-point fluid that, in turn, is used to drive a conventional turbine. To prevent flashing and scaling in the heat exchanger, a downhole pump is being considered. This system probably

will have about the same thermal efficiency as the flash system for brine deposits at 260 to 315°C. Better downhole pumps and more corrosion- and scale-resistant heat exchangers are necessary before this system can be applied commercially to high-salinity waters.

Since past research has shown the problems of scale deposition to be severe, LLL proposes an entirely different approach—the Total Flow System (Fig. 37; Austin, *et al.*, 1973). In principle, this system consists of passing the total, hot, wellhead, brine-steam mixture directly through a mixed-phase expander to drive a turbine and electrical alternator. One scheme is to expand the wellhead mixture through converging-diverging nozzles to convert fluid enthalpy into kinetic energy in the form of high-velocity jets, which can be used to drive a specially designed impulse turbine similar to hydroelectric

turbines. The important feature is that all pressure and temperature drop in the fluid will take place in the nozzles, allowing the turbine to operate at low temperature and pressure.

This concept offers the potential advantages of inherent mechanical and thermodynamic simplicity, high conversion efficiencies, and design flexibility, which facilitates scale control and permits a wider choice of materials. Also, by confining pressure and temperature drop to a single component (the nozzles) there is greater opportunity for brine treatment

and scale control. Successful development of this concept could yield about 60% more power output per wellhead product (i. e., an increase in overall thermal efficiency from 11% to about 18%). This is especially significant when applied to lower-temperature resources, in that fewer wells will be needed to produce the same electrical output. Because of this and the inherent mechanical simplicity of the system, capital investments and power costs should be lower than either flash-steam or binary-cycle processes.

Fig. 38. Lawrence Livermore Laboratory field material test chamber, well Sinclair #4, Imperial Valley, California.

Fig. 39. Lawrence Livermore Laboratory turbine test chamber, Livermore California.

In summary, the aim of the LLL geothermal program is to develop the technology necessary to convert high-temperature, high-salinity geothermal brines into useable electrical power.

Toward this end, major emphasis is on:

- Designing, testing, and evaluating energy-conversion concepts
- Analyzing brine chemistry and developing effective scale- and corrosion-control methods
- Identifying brine-tolerant materials and fabrication methods for geothermal applications
- Building and operating a small-scale geothermal power plant based on the Total Flow process.

The goal is to resolve most of the key technical issues by mid 1977 so that design of a 10-MWe demonstration system can begin that year. Completion and test of this pilot facility is scheduled for 1979. To date, progress in the LLL program includes: development of analytical methods for well flow characterization, initial studies of turbine configurations, construction of bench-model turbines, completion of a preliminary evaluation of the Salton Trough resource, analyses of typical scales taken from wells in the area, construction and operation of a field test chamber for preliminary material-screening and nozzle-performance tests (Fig. 38), and completion of a laboratory

Fig. 40. Lawrence Livermore Laboratory turbine test facility, Livermore, California.

for testing the performance of turbines, nozzles, and other components (Figs. 39 and 40). The performance-test facility consists of a water heater capable of producing about 0.7 kg/sec of 300°C saturated water at 1400 psia. By throttling to lower pressure, a range of thermodynamic characteristics typical of wellhead condi-

tions can be reproduced. The facility also includes a cooling system to allow turbine operation at backpressures as low as 12 kPa. The facility will be used for laboratory testing and performance evaluation of Total Flow concepts up to 100-kw capacity, prior to field testing.

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