PRELIMINARY STUDY OF THE URANIUM POTENTIAL OF THE TRIASSIC SANFORD BASIN AND COLON CROSS STRUCTURE, NORTH CAROLINA

BENDIX FIELD ENGINEERING CORPORATION
Grand Junction Operations
Grand Junction, Colorado 81501

January 1978

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SUMMARY

A preliminary geologic investigation was conducted to determine if the Triassic sedimentary rocks of the Sanford basin and Colon cross structure in North Carolina are favorable hosts for uranium deposits. Rocks of the adjacent Carolina slate belt were also examined as a potential source of uranium.

On the basis of favorability criteria for sandstone-type uranium deposits, and geologic and geophysical investigations of the study area, the most favorable sites for further investigation are (1) at the contacts between the Pekin and Cumnock and between the Pekin and Sanford Formations near the Colon cross structure and (2) at the base of the Jonesboro fault, which lies below the Sanford Formation, northwest of Sanford.

The highly weathered granites southeast of the Jonesboro fault were a source of the detritus deposited on the cross structure and may have been a primary source of uranium. Uranium leached from the coarse sediment (Pekin Formation) of the cross structure may have been transported downdip and may have been precipitated by the carbonaceous shales of the Cumnock Formation on the western side of the cross structure or at the Pekin-Sanford contact to the east.

The Jonesboro fault may provide an impermeable barrier to ground-water migration in the metamorphosed basement rocks below the Triassic sediments. Such a barrier would constitute a favorable site for the precipitation and retention of uranium.

Scintillometer surveys and laboratory analyses indicate no anomalous surface radioactivity in the study area. However, deep surface weathering may have caused the uranium to be leached from the exposed rocks and redeposited at depth.

Subsurface information was limited to rock cores from two test holes in the Sanford and Cumnock Formations in the northwestern part of the basin and to geophysical logs from a test hole to basement in the central part. A scintillometer scan of the core indicated no anomalous radioactivity, but examination of the gamma-ray logs revealed a slight increase in radioactivity in pyritic siltstones near basement in the Pekin Formation.

Geologic investigations show that conditions which have proven favorable for deposition of uranium in other areas are present in the Triassic rocks of the Sanford basin and Colon cross structure. However, because of deep surface weathering, further subsurface studies are necessary to confirm the favorability of the rocks as hosts for uranium.
INTRODUCTION

PURPOSE AND SCOPE

Many of the characteristics considered favorable for fluvial sandstone-type uranium deposits have been observed in rocks within the Triassic basins of North Carolina (Dennison and Wheeler, 1972, p. 9-14, 171-176). This preliminary study was designed to determine whether the sandstone formations in the 500-sq-mi Sanford basin and Colon cross structure in North Carolina are favorable as host rocks for uranium deposition, and to examine the possible source of uranium. The study was conducted by Bendix Field Engineering Corporation for the Grand Junction Office of the U.S. Energy Research and Development Administration, which became part of the Department of Energy on October 1, 1977.

The study involved 3 weeks of field work during October and November 1976. Several traverses were conducted across the basin, the cross structure, and the areas adjacent to the Triassic rocks to measure surface radioactivity and to collect samples. Radiometric measurements were obtained at 150 locations, and 42 samples were collected for laboratory analyses. Rock cores from coal test holes and geophysical logs from an oil test hole were examined for anomalous radioactivity. The sample locations and drill-hole locations are shown in Figure 1.

LOCATION

The Sanford basin and Colon cross structure constitute the southern half of the Deep River basin in central North Carolina. The study area is approximately 35 mi long by 15 mi wide and occupies parts of Lee, Moore, and Chatham Counties.

PREVIOUS WORK

Geologic investigations of the Triassic Deep River basin were conducted by Olmsted (1820), Emmons (1856), Campbell and Kimball (1923), Reinemund (1955), Mann and Zablocki (1961), Conley (1962), and Ackermann and others (1976).

Rock cores from coal tests conducted in Chatham County for the U.S. Bureau of Mines (Reinemund, 1955, p. 131-138) were made available by the North Carolina State Division of Mineral Resources. The same agency also provided geophysical logs from an oil test drilled in Lee County for the Chevron Oil Company during September and October 1964.

Stream-sediment and ground-water samples from the Sanford basin area have been analyzed for uranium by the Savannah River Laboratory. Stream-sediment samples collected from Moore County ranged from 0 to 500 ppm uranium (Savannah River Lab., 1976, p. 29). Precise sample locations were not available.

An airborne radiometric study of the Triassic basins in North Carolina was conducted in 1975 by Geodata International Incorporated under contract to the U.S. Energy Research and Development Administration.
Figure 1. Sample location map of the Triassic Sanford basin and Colon cross structure, North Carolina (after Reinemund, 1955).
PROCEDURES

FIELD PROCEDURES

Scintillometer scans were conducted across the study area and on rock cores from coal tests. Rock and soil samples were collected from outcrops in fields, at highway and railroad cuts, in rock quarries, in clay and gravel pits, and at excavation sites. Several contacts, unconformities, and fault zones, reported by Reinemund (1955) and Conley (1962), were investigated and sampled.

LABORATORY PROCEDURES

Analyses of rock and soil samples were performed at the analytical and petrographic laboratory of the Grand Junction Office of the Energy Research and Development Administration operated by Bendix Field Engineering Corporation. Gamma-ray spectroscopy was used to determine equivalent amounts of potassium, uranium, and thorium; fluorometric analysis was used to determine amounts of chemical uranium (Table 1). The techniques used in these studies were described by Latimer and others (1970, p. 138-141). Petrographic examinations consisted of modal analyses and mineral descriptions of thin sections.

GEOLOGY

STRATIGRAPHY

The sedimentary rocks in the Sanford basin and Colon cross structure are part of the Newark Group (Reinemund, 1955, p. 26). The three sedimentary units described by Campbell and Kimball (1923, p. 20) are the Pekin, Cumnock, and Sanford Formations (Figs. 2, 3; Pl. 1). The Pekin and Sanford Formations consist of red, brown, or purple arkosic and argillaceous rocks that grade from conglomerate or fanglomerate to claystone. Subordinate amounts of gray to brown arkosic sandstone and conglomerate are exposed throughout the study area. The Cumnock consists of gray and black arkosic and argillaceous rocks that grade from sandstone to claystone and shale. Two coal seams are also present in the Cumnock. The three units may be distinguished by differences in color and gross lithology, but the contacts are often gradational and difficult to locate (Reinemund, 1955, p. 39-53). Petrographic analyses of rock samples indicate mineral compositions and rock types consistent with those reported by Reinemund (1955, p. 39-52).

The basal conglomerate and sandstone of the Pekin Formation were derived from the pre-Triassic slate belt rocks, west of the Deep River basin. However, most of the sediments in the basin and cross structure were eroded from the highland southeast of the basin and were carried westward by streams. The coarser sediments of the Pekin and Sanford Formations were primarily deposited in channels and alluvial fans. The finer sediments were deposited on flood plains (Reinemund, 1955, p. 52-53). Reinemund (1955, p. 53) suggested that these sediments were derived from red soils of weathered pre-Triassic metamorphic
TABLE 1. LABORATORY ANALYSES OF ROCK AND SOIL SAMPLES FROM THE TRIASSIC SANFORD BASIN, COLON CROSS STRUCTURE, AND ADJACENT SLATE BELT

<table>
<thead>
<tr>
<th>Sample no.</th>
<th>Unit</th>
<th>U₃O₅ (ppm)</th>
<th>eU (ppm)</th>
<th>Th (ppm)</th>
<th>K (%)</th>
<th>Sample descriptions</th>
</tr>
</thead>
<tbody>
<tr>
<td>S-1</td>
<td>Sanford</td>
<td>1</td>
<td>1.7</td>
<td>7.5</td>
<td>0.9</td>
<td>Mottled red, gray-green sandstone; conglomerate</td>
</tr>
<tr>
<td>S-2</td>
<td>Sanford</td>
<td>2</td>
<td>1.7</td>
<td>9.2</td>
<td>2.0</td>
<td>Dark-red sandy siltstone</td>
</tr>
<tr>
<td>S-3</td>
<td>Sanford</td>
<td>1</td>
<td>0.5</td>
<td>7.8</td>
<td>1.1</td>
<td>Gray shaly siltstone</td>
</tr>
<tr>
<td>S-4</td>
<td>Cumnock</td>
<td>1</td>
<td>1.8</td>
<td>6.3</td>
<td>1.0</td>
<td>Gray sandy siltstone</td>
</tr>
<tr>
<td>S-5</td>
<td>Pre-Triassic</td>
<td>3</td>
<td>2.9</td>
<td>12.2</td>
<td>1.8</td>
<td>Orange-red soil</td>
</tr>
<tr>
<td>S-6</td>
<td>Pekin</td>
<td>1</td>
<td>1.1</td>
<td>3.3</td>
<td>0.6</td>
<td>Dark-red sandstone and conglomerate</td>
</tr>
<tr>
<td>S-7</td>
<td>Sanford</td>
<td>2</td>
<td>1.9</td>
<td>11.8</td>
<td>1.4</td>
<td>Orange-red soil</td>
</tr>
<tr>
<td>S-8</td>
<td>Tuscaloosa</td>
<td>3</td>
<td>1.5</td>
<td>5.8</td>
<td>1.6</td>
<td>White-gray sandstone</td>
</tr>
<tr>
<td>S-9</td>
<td>Pre-Triassic</td>
<td>2</td>
<td>1.4</td>
<td>7.1</td>
<td>2.8</td>
<td>Gray tuffaceous argillite</td>
</tr>
<tr>
<td>S-10</td>
<td>Pre-Triassic</td>
<td>1</td>
<td>0.5</td>
<td>0.6</td>
<td>0.0</td>
<td>Smokey vein quartz</td>
</tr>
<tr>
<td>S-11</td>
<td>Pekin</td>
<td>3</td>
<td>1.8</td>
<td>8.9</td>
<td>1.5</td>
<td>Dark-red sandy claystone</td>
</tr>
<tr>
<td>S-12</td>
<td>Pekin</td>
<td>2</td>
<td>2.0</td>
<td>8.6</td>
<td>1.6</td>
<td>Orange-tan sandy claystone</td>
</tr>
<tr>
<td>S-13</td>
<td>Pekin</td>
<td>2</td>
<td>0.7</td>
<td>3.8</td>
<td>0.5</td>
<td>Pebble conglomerate; red sandy matrix</td>
</tr>
<tr>
<td>S-14</td>
<td>Pekin</td>
<td>1</td>
<td>1.4</td>
<td>5.2</td>
<td>1.1</td>
<td>Red silty sandstone</td>
</tr>
<tr>
<td>S-15</td>
<td>Diabase dike</td>
<td>1</td>
<td>1.7</td>
<td>7.6</td>
<td>1.0</td>
<td>Tan-gray diabase</td>
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<tr>
<td>S-16</td>
<td>Pekin</td>
<td>2</td>
<td>0.8</td>
<td>7.4</td>
<td>0.6</td>
<td>Gray sandy claystone</td>
</tr>
<tr>
<td>S-17</td>
<td>Diabase dike</td>
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<td>2.1</td>
<td>5.1</td>
<td>0.5</td>
<td>Tan diabase</td>
</tr>
<tr>
<td>S-18</td>
<td>Pekin</td>
<td>2</td>
<td>1.3</td>
<td>6.4</td>
<td>1.1</td>
<td>Dark-red sandstone</td>
</tr>
<tr>
<td>S-19</td>
<td>Pekin</td>
<td>1</td>
<td>2.1</td>
<td>7.3</td>
<td>1.1</td>
<td>Gray bleached sandstone</td>
</tr>
<tr>
<td>S-20</td>
<td>Pre-Triassic</td>
<td>1</td>
<td>1.2</td>
<td>9.6</td>
<td>2.6</td>
<td>Gray argillite</td>
</tr>
<tr>
<td>S-21</td>
<td>Diabase dike</td>
<td>1</td>
<td>0.6</td>
<td>0.5</td>
<td>0.3</td>
<td>Dark gray-green diabase</td>
</tr>
<tr>
<td>S-22</td>
<td>Cumnock</td>
<td>2</td>
<td>1.2</td>
<td>7.3</td>
<td>1.1</td>
<td>Tan siltstone</td>
</tr>
<tr>
<td>S-23</td>
<td>Tuscaloosa</td>
<td>4</td>
<td>3.3</td>
<td>17.2</td>
<td>0.1</td>
<td>Tan sandstone</td>
</tr>
<tr>
<td>S-24</td>
<td>Pekin</td>
<td>1</td>
<td>1.1</td>
<td>4.2</td>
<td>1.4</td>
<td>Metagraywacke</td>
</tr>
<tr>
<td>S-25</td>
<td>Sanford</td>
<td>1</td>
<td>0.6</td>
<td>4.6</td>
<td>1.9</td>
<td>Fanglomerate</td>
</tr>
<tr>
<td>S-26</td>
<td>Granite intrusive</td>
<td>2</td>
<td>2.4</td>
<td>6.9</td>
<td>0.0</td>
<td>Gray weathered granite</td>
</tr>
<tr>
<td>S-27</td>
<td>Pre-Triassic</td>
<td>2</td>
<td>1.4</td>
<td>9.6</td>
<td>2.7</td>
<td>Rhyolite</td>
</tr>
<tr>
<td>S-28</td>
<td>Sanford</td>
<td>1</td>
<td>1.8</td>
<td>8.8</td>
<td>1.6</td>
<td>Purple-gray claystone; siltstone</td>
</tr>
<tr>
<td>S-29</td>
<td>Pre-Triassic</td>
<td>2</td>
<td>1.9</td>
<td>7.6</td>
<td>2.5</td>
<td>Weathered granite</td>
</tr>
</tbody>
</table>
TABLE 1. (continued)

<table>
<thead>
<tr>
<th>Sample no.</th>
<th>Unit</th>
<th>$U_3O_8$ (ppm)</th>
<th>eU (ppm)</th>
<th>Th (ppm)</th>
<th>K (%)</th>
<th>Sample descriptions</th>
</tr>
</thead>
<tbody>
<tr>
<td>S-30</td>
<td>Sanford</td>
<td>1</td>
<td>0.2</td>
<td>5.9</td>
<td>2.3</td>
<td>Gray silty claystone</td>
</tr>
<tr>
<td>S-31</td>
<td>Sanford</td>
<td>2</td>
<td>0.4</td>
<td>5.9</td>
<td>1.4</td>
<td>Gray weathered claystone</td>
</tr>
<tr>
<td>S-32</td>
<td>Sanford</td>
<td>1</td>
<td>0.6</td>
<td>5.9</td>
<td>1.1</td>
<td>White clay stringers</td>
</tr>
<tr>
<td>S-33</td>
<td>Pre-Triassic/Sanford</td>
<td>1</td>
<td>0.7</td>
<td>3.8</td>
<td>1.0</td>
<td>Fault contact between S-32 and S-34</td>
</tr>
<tr>
<td>S-34</td>
<td>Pre-Triassic</td>
<td>3</td>
<td>1.7</td>
<td>8.1</td>
<td>3.2</td>
<td>Orange-brown granite saprolite</td>
</tr>
<tr>
<td>S-35</td>
<td>Diabase dike</td>
<td>&lt;1</td>
<td>--</td>
<td>1.0</td>
<td>0.4</td>
<td>Gray diabase</td>
</tr>
<tr>
<td>S-36</td>
<td>Sanford</td>
<td>2</td>
<td>0.7</td>
<td>8.3</td>
<td>1.5</td>
<td>Red silty claystone</td>
</tr>
<tr>
<td>S-37</td>
<td>Cumnock</td>
<td>1</td>
<td>2.0</td>
<td>7.2</td>
<td>0.8</td>
<td>Gray silty claystone</td>
</tr>
<tr>
<td>S-38</td>
<td>Tuscaloosa</td>
<td>1</td>
<td>2.6</td>
<td>5.1</td>
<td>1.1</td>
<td>Gray-pink sandstone</td>
</tr>
<tr>
<td>S-39</td>
<td>Pekin</td>
<td>2</td>
<td>2.2</td>
<td>7.9</td>
<td>2.0</td>
<td>Red silty mudstone</td>
</tr>
<tr>
<td>S-40</td>
<td>Tuscaloosa/Pekin</td>
<td>3</td>
<td>2.6</td>
<td>12.1</td>
<td>2.0</td>
<td>Unconformable contact between S-39 and S-41</td>
</tr>
<tr>
<td>S-41</td>
<td>Tuscaloosa</td>
<td>2</td>
<td>1.6</td>
<td>4.3</td>
<td>1.6</td>
<td>Orange-brown sandstone</td>
</tr>
<tr>
<td>S-42</td>
<td>Tuscaloosa</td>
<td>1</td>
<td>1.4</td>
<td>8.1</td>
<td>0.4</td>
<td>Gray sandy siltstone</td>
</tr>
</tbody>
</table>
Figure 2. Geologic and structural map of the Triassic Sanford basin and Colon cross structure, North Carolina (after Reinemund, 1955).
Figure 3. Generalized stratigraphic section and description of rock units in the Triassic Sanford Basin (from Reinemund, 1955).
rocks that were transported into the basin by streams. The red coloration of the Pekin and Sanford rocks indicates deposition in an oxidizing environment. In contrast, the gray rocks of the Cumnock were deposited in a swampy or lacustrine reducing environment.

Reinemund (1955, p. 27) stated that the thickness of Triassic rocks is 7,000 to 8,000 ft in the Sanford basin and 4,000 to 5,000 ft in the Colon cross structure. Ackermann and others (1976, p. 139), who estimated thicknesses in the basin and cross structure from resistivity data, agreed in general with Reinemund. On the basis of gravity studies, however, Mann and Zablocki (1961, p. 212-213) suggested a maximum of 6,100 ft of sediment in the Sanford basin and 2,000 ft in the area of the Colon cross structure.

The diabase dikes are generally dark gray to olive black and contain plagioclase, pyroxene, olivine, and opaque minerals. The grain size ranges from 3.5 mm to less than 0.5 mm (Reinemund, 1955, p. 59). Dikes, which are as much as 300 ft wide and 7 mi long, have intruded northwest-trending joints and cross faults. Because they displace the longitudinal faults, intrude the youngest Triassic rocks, and do not intrude the Cretaceous or younger rocks, Reinemund (1955, p. 61) believed that the dikes probably were formed during the final stages of faulting near the end of the Triassic Period.

STRUCTURE

The Sanford basin is a half-graben that contains sedimentary rocks and occupies a northeast-trending depression in the Slate Belt rocks of the Piedmont province (Figs. 1, 2; Pl. 1). The basin probably originated as a down-faulted block, and the present structure evolved through differential subsidence accompanied by drag-folding and upward tilting of the beds along the Jonesboro fault (Reinemund, 1955, p. 75-76). The rocks in the basin generally dip southeastward at an average angle of approximately 15°. The southeastern boundary of the basin is the Jonesboro fault, which is more than 100 mi long and has a vertical displacement of approximately 6,000 to 10,000 ft. A series of longitudinal faults nearly parallel to the Jonesboro fault are cut and offset by a series of transverse faults and diabase dikes.

The Colon cross structure is a northwest-trending anticlinal cross fold that interrupts the northeast-trending synclinal form and separates the Durham and the Sanford basins. The cross fold affects both the Triassic strata and the pre-Triassic basement rocks. Reinemund (1955, p. 75) suggested that the cross fold was formed by sedimentation and the simultaneous movement of the Jonesboro fault. Later, the cross fold was cut by a series of transverse faults and dikes to form the present structure.

RESULTS

SURFACE STUDY

Radiometric surveys in the Sanford basin, the Colon cross structure, and the adjacent Slate Belt rocks revealed no anomalous radioactivity. Laboratory analyses of rock and soil samples did not disclose anomalous uranium
mineralization. However, Overstreet (1970, p. 381) noted that deep surface weathering in the southeastern states precludes significant ore deposits at the surface. Uranium possibly has been leached from exposed rocks and redeposited downdip.

SUBSURFACE STUDY

Several coal test holes were drilled in Lee and Chatham Counties through the Sanford and Cumnock Formations near the northwestern border of the Sanford basin (Reinemund, 1955, p. 127-156). A scintillometer scan of the cores from two of these holes revealed no anomalous radioactivity.

Geophysical logs of an oil test drilled in Lee County for the Chevron Oil Company showed an increased level of radioactivity in a gray to black, pyritic siltstone in the Pekin Formation, approximately 35 ft thick, near the metamorphosed basement rocks. However, the core from this hole was not available for analysis.

CONCLUSIONS

On the basis of criteria for sandstone-type uranium deposits (Grutt, 1972) the Sanford basin and Colon cross structure contain rocks which are favorable as hosts for uranium deposition. Oxidized, arkosic, fluvial-deltaic sandstones, claystones, and shales, which intertongue with reduced sandstones and claystones, are predominant in the Pekin and Sanford Formations. The Cumnock Formation contains carbonaceous shales that formed in a lacustrine environment and that may have caused reduction and precipitation of uranium. Faults and diabase dikes, which can form impermeable barriers to mineralized solutions, occur throughout the basin and along its borders.

According to Reinemund (1955, p. 53), the anticlinal Colon cross structure was the primary locus of drainage for the Sanford basin throughout most of Triassic time. Coarse sediments were deposited on the cross structure and finer sediments were deposited in the basins. Pre-Triassic granitic bodies, east of the cross structure and adjacent to the Jonesboro fault, were a source of detrital material and were the most likely source of uranium for the basin (Fig. 2).

West of the cross structure, the black shales of the Cumnock, which are in contact with and downdip from the coarse sandstones of the Pekin (Fig. 2), create a favorable situation for uranium reduction and deposition. Downdip from and northeast of the cross structure, the diabase dikes and faults may have provided effective traps for emplacement of uranium deposits. According to Reinemund (1955, p. 40), an angular unconformity separates the Pekin from the overlying Sanford Formation along the eastern and southeastern limbs of the cross structure. Grutt (1972, p. 50) suggested that unconformities may play a significant role in uranium deposition. From gravity surveys in the study area, Mann and Zablocki (1961, p. 206-208) reported that 2,000 ft is the maximum thickness of the Triassic sediments parallel to the contact between the Pekin and Sanford Formations northeast of the Colon cross structure. The sediments on the southwestern side are approximately the same thickness.
The Jonesboro fault, which bounds the deepest part of the basin, may provide an impermeable trap for uranium moving near the metamorphosed basement rocks. Mann and Zablocki (1961, p. 208) state that the Sanford basin attains a maximum depth of approximately 5,800 ft northwest of Sanford. This area is downdip from both the western border and the Colon cross structure. The Chevron oil test showed abundant pyrite, which is a good reductant for uranium, in the Pekin Formation approximately 4 mi northwest of Sanford (Fig. 1).
BIBLIOGRAPHY


