

Competitive sorption of cadmium and lead in acid soils of Central Spain

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Summary

The bioavailability and ultimate fate of heavy metals in the environment are controlled by chemical sorption. To assess competitive sorption of Pb and Cd, batch equilibrium experiments (generating sorption isotherms) and kinetics sorption studies were performed using single and binary metal solutions in surface samples of four soils from central Spain. For comparisons between soils, as well as, single and binary metal solutions, soil chemical processes were characterized using the Langmuir equation, ionic strength, and an empirical power function for kinetic sorption. In addition, soil pH and clay mineralogy were used to explain observed sorption processes. Sorption isotherms were well described by the Langmuir equation and the sorption kinetics were well described by an empirical power function within the reaction times in this study. Soils with higher pH and clay content (characterized by having smectite) had the greatest sorption capacity as estimated by the maximum sorption parameter (Q) of the Langmuir equation. All soils exhibited greater sorption capacity for Pb than Cd and the presence of both metals reduced the tendency for either to be sorbed although Cd sorption was affected to a greater extent than that of Pb. The Langmuir binding strength parameter (k) was always greater for Pb than for Cd. However, these k values tended to increase as a result of the simultaneous presence of both metals, that may indicate competition for sorption sites promoting the retention of both metals on more specific sorption sites. The kinetic experiments showed that Pb sorption is initially faster than Cd sorption from both single and binary solutions although the simultaneous presence of both metals affected the sorption of Cd at short times while only a minor effect was observed on Pb. The estimated exponents of the kinetic function were in all cases smaller for Pb than for Cd, likely due to diffusion processes into micropores or interlayer space of the clay minerals which occurs more readily for Cd than Pb. Finally, the overall sorption processes of Pb and Cd in the smectitic soil with the highest

sorption capacity of the studied soils are slower than in the rest of the soils with a clay mineralogy dominated by kaolinite and illite, exhibiting these soils similar sorption rates. These results demonstrate a significant interaction between Pb and Cd sorption when both metals are present that depends on important soil properties such as the clay mineralogy.

Keywords:

Lead; Cadmium; Sorption isotherms; Sorption kinetics; Competitive sorption; Acid soils.

1. Introduction

Metals are natural constituents of soils. However, in the last decades, significant changes in the global budget of heavy metals at the earth's surface have occurred (Förstner, 1995). Industrial activities, fertilizer and sewage sludge applications as well as effluent disposal on land can result in significant input of heavy metals. This can lead to either substantial accumulation, in excess of the natural background, or leaching, potentially polluting surface or subsurface water bodies, or both.

Two of the most potentially toxic heavy metals are cadmium and lead. Classified as soluble and strongly hydrating cations (McBride, 1994), both metals are particularly toxic to higher animals, producing kidney and blood diseases among other health disorders. The term 'sorption' is used to describe the removal of metals in solution by the soil solid phase (Hooda and Alloway, 1994; Chen et al., 1997) including any retention mechanism that controls availability and mobility. Many studies have focused on the sorption of these metals on different soil materials and under different experimental conditions (Hooda and Alloway, 1998; Martínez and McBride, 1998; Sauvé et al., 2000; Appel and Ma, 2002; Krishnamurti and Naidu, 2003; Trivedi et al., 2003; Adhikari and Singh, 2003). Soil pH, other factors such as the presence of competing ligands, the ionic strength of the soil solution, and the simultaneous presence of competing metals are known to significantly affect sorption processes and leaching potential through a soil profile (Kookama and Naidu, 1998; Harter and Naidu, 2001).

However, despite the established significance of competitive sorption, and that sorption selectivity for a particular metal might result from its relative affinity for specific sites or its sorption on to sites unavailable to other metals (Benjamin and Leckie, 1981b), most soil-metal bonding information has been derived from studies conducted using single metal solutions. Also, while monoion sorption studies may

adequately predict sorption of strongly bonded ions, sorption of less strongly bonded ions is more likely to be affected by the presence of competing ions in solution (Harter, 1992). Such studies may have limited practical applications when used to explain sorption in soils containing competitive cations (Fontes et al., 2000). Greater progress has been made in studying competitive sorption reactions in pure minerals and organic compounds (Kinniburgh et al., 1976; Tiller et al., 1979; Benjamin and Leckie, 1981b; Elliott et al., 1986; Bereket et al., 1997; Pinheiro et al., 1999; Saha et al., 2002) than in more heterogeneous soil surfaces (Cavallaro and McBride, 1978; Murali and Aylmore, 1983; Harter, 1992; Mesquita and Viera e Silva, 2002). However, theoretical sorption models based on simple mineral or organic systems appear unlikely to provide the means for quantitative predictions in complex soils (Tiller et al., 1984). Metal characteristics such as the charge-to-radius ratio (Gomes et al., 2001) or metal-ion hydrolysis constants (Welp and Brümmer, 1999) sequences do not always explain metal bonding selectivity to heterogeneous soil systems.

Therefore, we designed an investigation to examine the effect of the simultaneous presence of Pb and Cd on sorption behavior to acidic A horizons of temperate soils from Spain. Specifically, we have undertaken batch equilibrium experiments to generate sorption isotherms and kinetic sorption studies using single and binary metal solutions in four soils characterized with different proportions of variable- and permanent-charge clay minerals. Results were characterized and compared for different metal solutions and soils using the Langmuir equation, ionic strength, the total metal retained in the soils, soils characteristics (including pH and clay content), and an empirical power function for kinetic sorption.

2. Materials and Methods

2.1. Soils

We collected bulk samples of the topsoils (0-15 cm depths) from four acidic soils. Three soils (S1, S2, and S3) were developed from Pliocene–Quaternary aged formations (raña) in Cáceres, Spain. They were classified as a Plinthic Palexerult (S1), Ultic Palexeralf (S2), and Arenic Pachic Palexerult (S3) (Soil Survey Staff, 1999). The fourth soil was developed on a hillslope in Madrid (Spain) and was classified as a Vertic Haploxerert (S4). To obtain a homogeneous sample of the top soils at each location, three approximately 3-kg samples from 2 m apart of the same horizon of each soil were combined prior to the experiments. All samples were air-dried, crushed and sieved through a 2-mm mesh prior to soil characterization and sorption studies.

2.2. General soil analyses

Soil pH was measured in deionized water (pH_w) and in 1 M KCl (pH_k) (in a 1:2.5 suspension), and organic carbon (OC) was determined by wet digestion (Walkley and Black, 1934). The exchangeable bases were extracted with 1 M NH_4OAc (pH 7) (Thomas, 1982), and the exchangeable aluminum (Al_k) was extracted with 1 M KCl (Barnhisel and Bertsch, 1982). The effective capacity of the exchange complex (ECEC) was calculated as the sum of Al_k and the amounts of Ca, Mg, Na and K extracted by 1 M NH_4OAc at pH 7 (Shuman, 1990). The supernatants from each extraction were separated by centrifuging at 6640 g for 20 minutes and stored in polyethylene containers at 4 °C prior to analysis. Determinations were made in triplicate. Blank extractions (without soil) were carried out for each set of analyses. The Ca, Mg, and Al contents were determined by ICP-AES on a Perkin Elmer OPTIMA 4300DV, K and Na by flame

emission spectroscopy on an Eppendorf ELEX 6361 instrument, and Al_K on a Perkin Elmer 403 atomic absorption spectrometer.

The mineralogical compositions of the total (≤ 2 mm) and clay (≤ 2 μ m) fractions were identified by X-ray powder diffraction on a Philips X'Pert diffractometer with graphite-monochromated CuK $_{\alpha}$ radiation. The XRD patterns were obtained from random powder mounts and various oriented aggregates of the Mg- and K- clay (air-dried, ethylene glycol-solvated, heated at 300 °C for 3 hours, and heated at 500 °C for 3 hours). We obtained semi-quantitative estimates of the minerals from random powder and oriented aggregated patterns, using intensity factors reported by Schultz (1964).

2.3. Sorption experiments

Both kinetic and isotherm experiments of Cd and Pb from their single and binary mixed solutions were undertaken using a batch equilibrium technique. Batch experiments were performed by adding 15 ml of single- (Pb or Cd) or binary-metal (Pb + Cd) solutions to duplicate 10-g soil samples in 50-ml polypropylene centrifuge tubes. Kinetic experiments were performed using four different initial metal concentrations. Each initial solution of both metals in single and binary solutions was prepared with similar total ionic strength (Table 1). Ionic strength (mM) was calculated by:

$$I = 1/2 \sum_i C_i Z_i^2 \quad (1)$$

where C_i is the concentration (mmol dm^{-3}) of the i th species, Z_i is its charge and \sum extends over all the ions in solution (Sparks, 1995). In the binary solutions, in order to achieve the same I value of the single solutions, the concentration of each metal was proportional to the molecular weight of the metals. All solutions were prepared from chloride metal salts in a 1 mM CaCl₂ background solution. Given the different metal sorption capacity of the soils and based on preliminary studies, we established similar

total initial ionic strength of all the solutions for S1 and S3. A different total initial ionic strength solution was used for S2 and S4 (Table 1). Theoretical calculations using MINTEQA2 (USEPA, 1997) indicated that all initial solutions were undersaturated with respect to each metal chloride and hydroxide at the pH of the solutions. The soil suspensions for each initial metal solution concentration were shaken on an end-over-end shaker (30 rpm) in a controlled room temperature ($25^{\circ}\text{C} \pm 2$) for 1, 5, 15, 30, 60, 180, 300, 900, and 1440 (24 h) minutes. The suspensions were centrifuged at 6640 g for 20 minutes and the supernatants removed by filtration (Whatman No. 42) before the pH of the solutions were determined and the solutions were analysed for Cd and Pb by ICP–AES.

The total amount of metal retained by the solid phases was obtained by

$$S = (C_o - C_t)V / W \quad (2)$$

where S is the amount of metal sorbed per unit mass of soil ($\mu\text{mol kg}^{-1}$), C_o and C_t are the metal concentration in the initial solution and after the reaction time in the filtrate, respectively ($\mu\text{mol dm}^{-3}$), V is the volume of solution added (dm^{-3}) and W is the air-dried mass of soil (g).

For each reaction time, sorption isotherms of the metals from their single and binary solutions were constructed using the data obtained from the kinetic experiments and the resulting sorption data obtained from two additional initial concentrations (Figure 1, Table 1). The corresponding sorption isotherms for each equilibration time were investigated by fitting the experimental data to the Langmuir isotherm given by:

$$S = \frac{QkC}{1 + kC} \quad (3)$$

where S is the amount of heavy metal sorbed by the soil solids ($\mu\text{mol kg}^{-1}$), C the equilibrium concentration in solution ($\mu\text{mol dm}^{-3}$), Q the maximum sorption ($\mu\text{mol kg}^{-1}$) and k the bonding energy coefficient (reciprocal $\mu\text{mol dm}^{-3}$) (Kinniburgh, 1986).

An empirical power function was fitted to experimental data from the kinetic sorption experiments (Kuo and Lotse, 1974; Aharoni and Sparks, 1991; Sparks, 1995) as:

$$S_t = kt^v \quad (4)$$

where S_t is the amount of metal retained at time t ($\mu\text{mol dm}^{-3}$), t is the reaction time (min), and k and v are constants and v is positive and less than unity (Sparks and Jardine, 1984; Chien and Clayton, 1980).

The optimal parameter values for both Langmuir and kinetic equations were determined by non-linear regression analysis, on the assumption of a constant relative error as the residuals revealed no systematic deviation. The goodness-of-fit for both the isotherms and kinetic equations was estimated by the coefficient of determination (R^2), the confidence intervals (95 %) of the estimated parameters (CI), and the standard error of the estimate (SE).

3. Results and discussion

3.1. Soil characteristics

All soils were moderately acidic in the surface horizons with differences between pH_w and pH_K close to or greater than 1. In S1 and S3, Al accounted for the 70 % and 64 % of the ECEC, respectively, while Ca was the dominant cation in S2 and S4 accounting for the 52 % and 69 % of the ECEC, respectively. Soils S1 and S3 had the lowest ECEC, pH_w and clay content while S2 and S4 were less acidic and showed higher ECEC and clay content. Soil S4, with the highest organic carbon and clay content had the greatest ECEC value ($20 \text{ cmol}_c \text{ kg}^{-1}$). Other soil properties are shown in Table 2.

In addition to the differences in the pH, organic carbon, and clay content, the four soils differed in the mineralogical composition of the clay fraction which also conditions their relative sorptive properties (Table 2). The clay fraction of S1 was dominated by kaolinite and to a lesser extent by illite. Soil S2 contained less kaolinite and more smectite and illite. Both S1 and S2 had similar proportions of phyllosilicates in the ≤ 2 mm soil fraction and of goethite and haematite in the clay fraction. The greater content of illite and the moderate quantity of smectite, provide S2 with greater exchange capacity than S1, as these minerals have larger net surface charge than kaolinite. Also, the presence of smectite provides S2 with permanent surface charge.

The clay fraction of soil S3 consisted predominantly of illite with less kaolinite than soils S1 and S2. Soil S3 had the lowest clay content (45 g kg^{-1}) and the smallest proportion of phyllosilicates in the ≤ 2 mm soil fraction. This soil also had a pH_w of 5 with low organic carbon (3 g kg^{-1}). As a result, S3 had the lowest exchange capacity. The clay fraction of soil S4 was dominated by well crystallized smectite and a sizable proportion of illite that provide the soil with permanent surface charge. In addition, this soil exhibited both the largest clay content (230 g kg^{-1}) and proportion of phyllosilicates in the ≤ 2 mm soil fraction which, along with the organic carbon content, justify its largest exchange capacity of the four soils. The presence of smectite as the dominant clay ensures high metal sorption capacity (Veeresh et al., 2003) as it provides the soil with high cation exchange capacity, an established factor regulating the sorption of heavy metals by soils (Kuo and Baker, 1980; Hooda and Alloway, 1998; Gomes et al., 2001; Appel and Ma, 2002).

3.2. Sorption isotherms.

Cadmium and Pb sorption data (1440 minutes reaction time) for both single and

binary initial solutions, were adequately described by the Langmuir equation with high R^2 and low values of SE (Figure 1, Table 3). Langmuir parameters Q and k were not correlated. Soils S1 and S3 had lower metal sorption capacity than S2 and S4 in terms of the estimated maximum sorption parameter Q (Q_{Pb} and Q_{Cd} for single metal solutions, and Q_{Pb}^* and Q_{Cd}^* , for binary metal solutions) (Table 3). This would be expected given the pH, clay content, and its mineralogical composition of the soils (Table 2). Soil S4 had the highest Q values, that would also be expected due to the higher organic matter and clay contents, as well as, the clay mineralogical composition (high proportion of smectite) (Table 2). Soil S3 contained 71 % sand, a clay fraction characterized by a low exchange capacity, low organic matter, and the lowest pH. Accordingly, this soil generally had the lowest Q values (except for Q_{Cd}^* in S1). All soils exhibited similar sorption patterns, with Q values for Pb higher than for Cd regardless of whether the metals were applied in single or binary solutions. With the exception of S3, all soils had ratios Q_{Pb}/Q_{Cd} (ranging from 1.2 to 1.8) lower than ratios Q_{Pb}^*/Q_{Cd}^* (ranging from 2.1 to 3.4). In the case of S3, both ratios were more similar than in the other soils (1.8 and 1.5 for single and binary solutions, respectively). This confirms the higher affinity of Pb than Cd for sorbent surfaces generally found in both pure soil components and heterogeneous soils (Kinniburgh et al., 1976; Elliott et al., 1986; Appel and Ma, 2002; Gomes et al., 2001; Adhikari and Singh, 2003; Fontes and Gomes, 2003).

The Q_i/Q_i^* ratios were generally greater than unity (except for in S3) suggesting that the simultaneous presence of both metals reduced sorption through competition for sorption sites in the solid phases. In addition, it was generally true that $Q_{Cd}/Q_{Cd}^* > Q_{Pb}/Q_{Pb}^*$ suggesting that Cd sorption was more affected by the simultaneous presence of a competing metal than Pb. This tendency of Pb to effectively compete for sorption sites

on different colloidal surfaces has been described in the presence of Cd (Fontes and Gomes, 2003; Rodríguez-Maroto et al., 2003), Cu (Christl and Kretzschmar, 1999), and other metals in multimetal solutions (Fontes et al., 2000; Trivedi et al., 2001; Saha et al., 2002). At low concentrations, no competition between Pb and other metals were observed in other cases (Benjamin and Leckie, 1981a; Saha et al., 2002). In the case of S3, Q_{Cd}/Q_{Cd}^* and Q_{Pb}/Q_{Pb}^* ratios were similar and smaller than unity.

The bonding energy coefficient (k_{Pb} and k_{Cd} for single metal solutions, and k_{Pb}^* and k_{Cd}^* for binary metal solutions) varied with soil type and metal solution, although all soils showed greater affinity for Pb than for Cd as $k_{Pb} > k_{Cd}$ and $k_{Pb}^* > k_{Cd}^*$ (Table 3). Adhikari and Singh (2003) found similar results for single metal solutions, Rodríguez-Maroto et al. (2003) for both single and binary solutions, and this result also agrees with the generally accepted metal affinity series for soils and soil components (Elliott et al., 1986). However, in contrast to those authors but in agreement with (Mesquita and Viera e Silva, 2002) for competitive sorption of Cu and Zn, our study found $k_i \leq k_i^*$ in all soils except for in soil S3. While binding strength, or affinity constant (k), estimates made from sorption isotherms should only be considered qualitatively (Harter, 1984; Sparks, 1995), they have been related to the free energy change of adsorption of different species (Van Riemsdijk et al., 1985). Higher k values have been related to specifically sorbed metals at high energy surfaces with low dissociation constants. Alternatively, lower k values appear to be related to sorption at low energy surfaces with high dissociation constants (Ma and Rao, 1997; Adhikari and Singh, 2003). The bimetal isotherm k values in all soils except for S3, may indicate that competition for sorption sites promotes the retention of both metals on more specific sorption positions. As a result, although maximum sorption coefficient (Q_i) decreases, the metals are held more strongly. The irregular sorptive behavior of S3 in this regard could be explained by the

high metal load relative to its low sorption capacity as measured by the ECEC. Thus, the estimated k_i^* values decreased as a consequence of the increased sorption levels (McBride, 1999).

In order to examine the role of soil pH and hydrolysis, the sorption data were also plotted against the pH of the filtrated solutions after the equilibration time of 24 h and fitted to an exponential growth function linearized as

$$\ln S = mpH + b \quad (5)$$

where S is the amount of heavy metal sorbed by the soil solids ($\mu\text{mol kg}^{-1}$) and, m and b the slope and intercept, respectively. Similar to Mesquita and Vieira e Silva (2002), Fontes and Gomes, (2003), and Rodriguez-Maroto et al., (2003), the pH of the filtrated solutions consistently decreased with the sorption level (S_i) yielding negative values of the slope in equation 5 (Table 4). This has been attributed to metal hydrolysis and the displacement of exchangeable H^+ by the metal cations. However, in the single metal isotherms the slopes (Eq. 5) for Pb in the soils in this study were less negative than for Cd (Table 5). This may be due to the greater dependency for Cd retention on electrostatic interactions with exchange sites than Pb, where sorption is more dependent on the covalent interactions with the mineral structures (McBride, 1989; Appel and Ma, 2002).

In contrast, in the binary solutions, the slope for Pb tended to be slightly more negative than for Cd. In this case, the effect of the sorption of each metal on the final pH at equilibrium is difficult to assess in these experimental conditions. However, it has been stated that strongly adsorbing cations compete more efficiently with protons in acquiring their position in the electronic clouds of O atoms than do the weakly adsorbing cations (Abd-Elfattah and Wada, 1981). On the other hand, the value of the slopes for S1 and S3 were always less negative than those for S2 and S4 in both Pb and

Cd from single metal solutions. This difference may be the result of the higher pH and the greater clay content of S2 and S4. This last difference could induce a lowering of the pH attributed to enhanced hydrolysis of the metals to a greater extent than in S1 and S3 (McBride, 1989). However, this tendency was reversed for the binary metal solutions and we are unable to offer any explanation for this result.

3.3. Sorption kinetics

The kinetics of Pb and Cd sorption at all initial concentrations and from both single and binary solutions showed a two stage time-dependent behavior with an initially rapid reaction followed by a much slower phase, although some differences were observed between the metals, solutions, and the soils (Figure 2). Lead was initially more rapidly sorbed than Cd in all soils and from both single and binary solutions (Rodríguez-Maroto et al., 2003). For example, from the single solution of the lowest concentration and at an equilibration time of 15 min, more than 99 % of the initial Pb concentration was sorbed in S1, S2 and S4, and 96 % in S3. In contrast, under similar conditions, Cd sorption reached the 77 % of the initial concentration in S1, S2, and S3, and 91 % in S4. While these percentages did not vary for Pb from binary solutions, Cd sorption at 15 min increased, resulting in more than 87 % of the initial concentration in all soils. Overall, this initial rapid reaction that both metals underwent in single and binary solutions is characteristic of heavy metal sorption on pure components and soils and has been attributed to chemisorption on phyllosilicates (Eick et al., 2001), adsorption on high affinity surface sites (Glover et al., 2002), or on sites with higher bonding strength with the metal (McBride, 1999). Consequently, the increment in the initial Cd sorption rate in binary solutions could indicate that the competitive Pb sorption forces Cd retention on sorption sites with greater affinity or more specific for

this metal. However, the sorption mechanisms responsible for the slow reaction phase are not well understood (Glover et al., 2002) although it has been attributed to diffusion, precipitation, and/or sorption reactions on sites with higher activation energy than the fast sorption sites (Strawn and Sparks, 2000). Thus, the apparent rate coefficient of metal sorption reactions are composed of various chemical and diffusive reactions, difficult to differentiate in complex soil matrices from time dependent data without spectroscopic evidence (Glover et al., 2002).

Consequently, a fractional power function (Eq. 4) (Aharoni and Sparks, 1991) was used to compare the overall sorption kinetics of the metals in single and binary solutions. This equation is empirical and therefore its use does not support mechanistic information but simply provide a consistent method to compare experimental results. In general, Equation 4 adequately described the rate of metal sorption within the time ranges used in these experiments and from both single and binary solutions (given the high R^2 and low SE values).

The estimated exponents of Equation 4 (ν) for Pb and Cd sorption from single and binary solutions are listed in Table 5. These values could be related to the empirical rate coefficients of the overall sorption processes over the entire reaction time range. As can be seen in Table 4, in all soils, the simultaneous presence of the competing metal did not affect the estimated apparent sorption rate ($\nu_i \approx \nu_i^*$) at any initial concentration. This could indicate that among the different sorption processes that take place during the metal interaction with the soil components, the rate limiting factor, although difficult to identify, may not be affected by the presence of the competing metal. On the other hand, although due to the strong affinity of the soil for Pb (Strawn and Sparks, 2000) its sorption was apparently more rapid at short reaction times than Cd sorption, estimated ν values for Pb were consistently smaller than for Cd in all soils from both

single and binary solutions over the entire reaction time range ($v_{Cd} > v_{Pb}$ and $v_{Cd}^* > v_{Pb}^*$). This could be related to the greater tendency of Pb to be adsorbed as a hydrolyzed species than Cd (Glover et al., 2002), which limits the rate of Pb diffusion into micropores created by structural defects of the clay particles (Glover et al., 2002) or into the narrow interlayer space of 2:1 clay minerals (McBride, 1994). Finally, estimated apparent sorption rates for each metal from single and binary solutions are similar in soils S1, S2, and S3 and about one order of magnitude lower in S4 than in the rest of the soils. As described above, the mineralogical composition of soils S1, S2, and S3 is dominated by kaolinite and illite, whereas, the clay fraction of soil S4 is dominated by smectite. Metal sorption on kaolinite and illite does not differ much (Lackovic et al., 2004) and it is known to be a rapid reaction since their exchange capacities are mainly due to external surface and edge sites readily accessible to cation exchange (Jardine and Sparks, 1984). In fact, cation exchange on clays without narrow interlayer space such as kaolinite appears to be instantaneous in comparison to exchange on smectite which can be related to its freely expanding interlayer space (Jardine and Sparks, 1984) and limited by the rate of cation diffusion through this region (McBride, 1994).

This different sorptive behavior of the soils as a function of the clay mineralogical composition can be observed when sorption isotherms are constructed for each equilibration time (1 to 1440 minutes). The estimated Q values from fitting the Langmuir equation to all of the isotherms were plotted against the equilibration time for each soil and metal in both single and binary solutions. An empirical power function was found to adequately describe the resulting $Q(t)$ plots with R^2 values ranging from 0.83 to 0.99. The exponents of each function $Q(t)$ along with the corresponding R^2 values, confidence intervals of the estimated exponents, and the standard error of the estimate are shown in Table 6. The average value of the exponents found for S1, S2,

and S3, were 0.03 ± 0.015 (standard deviation) and 0.035 ± 0.003 for Pb, and 0.042 ± 0.017 and 0.040 ± 0.019 for Cd, in single and binary solutions, respectively. These values are lower than those found for soil S4. Based on this, it seems that the time-evolution of the Q value can also be related to clay mineralogical composition to a greater extent than to whether Pb and Cd are applied in single or binary solutions.

4. Conclusions

We performed a detailed investigation of competitive sorption processes between Pb and Cd metals using batch sorption isotherms and kinetics sorption studies for single and binary metal solutions in four soils. Sorption isotherms for Pb and Cd in single and binary solutions of similar total ionic strength were adequately described by Langmuir equation. The sorption capacity of the soils for lead, as measured by the estimated Q parameter from Langmuir equation, is greater than for cadmium. The co-existence of both metals reduces their tendency to be sorbed on the soil solid phases affecting, to a greater extent, the sorption capacity of Cd than Pb. Soils S2 and S4, with higher pH and clay content characterized by having sizable proportions of smectite, had the greatest metal sorption capacity as the presence of this clay mineral provides the soil with a large cation exchange capacity. On the other hand, in agreement with the metal affinity series, the binding strength parameter k was always greater for lead than for cadmium. However, in all soils (except for the S3) the simultaneous presence of both metals increased their corresponding k values indicating that competition for sorption sites could promote the retention of both metals on more specific sorption positions, although the amount of metal retained in the soil decreased. Therefore, results from an assessment of the potential bioavailability and toxicity of lead and cadmium might be different whether the experiments are performed using single or binary solutions.

The kinetics of Pb and Cd sorption from both single and binary solutions followed a two stage time-dependent behavior with an initially rapid reaction followed by a much slower stage. This sorption kinetic could be well described by an empirical power function within the reaction time ranges used in this study. The estimated apparent sorption rates of the metals from single and binary solutions were similar. This result could indicate that the rate limiting metal sorption stage in these soils for each metal is not significantly affected by the simultaneous presence of both species. On the other hand, Pb was initially more rapidly sorbed than Cd in all soils and from both single and binary solutions. However, the estimated exponents were in all cases smaller for Pb than for Cd, likely due to diffusion processes into micropores or interlayer clay spaces for which Cd could exhibit greater ease than Pb. Finally, the mineralogical composition of the clay fraction of these soils determined the empirical metal sorption rate within the time ranges employed in this study. The overall sorption processes of Pb and Cd in the smectitic soil S4, with the highest sorption capacity of the soils are slower than in the rest of the soils with a clay mineralogy dominated by kaolinite and illite.

From these results, it can be concluded that the sorption behavior of Pb and Cd in the moderate acidic soils is significantly affected by the simultaneous presence of both metals. Thus, the competitive sorption should be considered to correctly assess their potential bioavailability, toxicity, and leachability in soils.

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Figure captions

Figure 1.

Lead and cadmium sorption isotherms from both single (open symbols) and binary (closed symbols). Solid lines are the best fits to the Langmuir equation.

Figure 2.

Sorption kinetics of lead and cadmium from single and binary solutions of varying initial concentrations in S1 and S2.

Table 1
Initial metal concentrations and total ionic strength used in the sorption isotherms and kinetic sorption experiments

	S1		S2		S3		S4	
	C ₀ mM	Total ^a I ₀ mM	C ₀ mM	Total I ₀ mM	C ₀ mM	Total I ₀ mM	C ₀ mM	Total I ₀ mM
Pb	0.02*	3.07	0.89*	5.68	0.02*	3.07	4.15	15.45
	0.07*	3.20	1.85*	8.54	0.07*	3.20	5.55	19.65
	0.14	3.43	4.15	15.45	0.14	3.43	7.14	24.43
	0.35	4.04	5.55	19.65	0.35	4.04	11.92	38.76
	0.89	5.68	7.14	24.43	0.89	5.68	15.84*	50.51
	1.85	8.54	10.14	33.41	1.85	8.54	18.54*	58.62
Cd	0.04*	3.11	0.04*	3.11	0.04*	3.11	1.90*	8.69
	0.09*	3.26	0.06*	3.17	0.09*	3.26	2.74*	11.21
	0.19	3.56	3.58	13.75	0.19	3.56	3.61	13.83
	0.33	3.98	5.33	19.00	0.36	4.07	6.28	21.83
	0.89	5.66	6.82	23.45	0.89	5.66	8.22	27.65
	1.78	8.33	10.84	35.52	1.78	8.33	9.64	31.92
Pb(+Cd)	0.11	3.51	0.62*	5.84	0.11	3.51	2.83	16.36
	0.24	4.10	1.17*	8.47	0.24	4.10	3.57	20.29
	0.62	5.84	2.83	16.17	0.62	5.98	5.12	28.15
	1.17	8.47	3.76	20.86	1.17	8.71	7.48	39.96
	2.83*	16.36	4.34	24.33	2.56*	15.55	9.51*	50.05
	3.57*	20.29	7.14	39.94	3.57*	24.65	11.26*	58.55
Cd(+Pb)	0.06	3.51	0.33*	5.84	0.06	3.51	1.62	16.36
	0.13	4.10	0.65*	8.47	0.13	4.10	2.19	20.29
	0.33	5.84	1.56	16.17	0.37	5.98	3.27	28.15
	0.65	8.47	2.19	20.86	0.74	8.71	4.84	39.96
	1.62*	16.36	2.76	24.33	1.62*	15.55	6.18*	50.05
	2.19*	20.29	5.17	39.94	3.64*	24.65	7.26*	58.55

^a Total ionic strength of the solution as calculated by eq. 1

* Initial concentrations followed by * were used as additional solutions for the sorption isotherm experiments.

Table 2.
Physical and chemical properties of the experimental soils

Soil No.	pH _w ^a	pH _k ^b	OC	Ca	Mg	Na	K	Al	ECEC	Sand	Silt	Clay
			g kg ⁻¹	cmol _c kg ⁻¹						g kg ⁻¹		
S1 ^c	5.2 (0.02)	4.2 (0.03)	17 (1.0)	0.32 (0.02)	0.08 (0.00)	0.04 (0.00)	0.09 (0.00)	1.26 (0.08)	1.79 (0.11)	675 (38)	250 (15)	75 (5)
S2	6.1 (0.05)	4.5 (0.02)	12 (0.6)	1.58 (0.04)	0.49 (0.01)	0.07 (0.01)	0.68 (0.12)	0.19 (0.01)	3.01 (0.18)	643 (35)	201 (13)	156 (12)
S3	5.0 (0.02)	3.8 (0.01)	3 (0.1)	0.06 (0.00)	0.20 (0.02)	0.02 (0.00)	0.11 (0.01)	0.7 (0.03)	1.09 (0.05)	710 (42)	245 (16)	45 (6)
S4	5.8 (0.03)	4.3 (0.02)	61 (6.5)	14.10 (0.58)	5.22 (0.50)	0.21 (0.03)	0.39 (0.02)	0.56 (0.05)	20.5 (1.25)	560 (25)	210 (8)	230 (19)

^apH_w, pH measured in deionized water

^bpH_k, pH measured in 1 M KCl

^cMean values and standard deviation between parenthesis (n=3)

Semi-quantitative mineralogical composition (relative % between samples) of the soils.

Fraction	Soil No	Q	F _{Ca-Na}	F _K	G	H	Ph	S	V	I	K
≤ 2 mm	S1	74	tr	tr	4	3	19	-	-	-	-
	S2	80	tr	tr	tr	nd	20	-	-	-	-
	S3	52	8	31	nd	nd	9	-	-	-	-
	S4	40	18	9	nd	nd	33	-	-	-	-
≤ 2 μm	S1	12	nd	nd	6	6	76	nd	8	11	57
	S2	22	nd	nd	7	3	68	14	nd	22	32
	S3	13	5	4	nd	nd	78	nd	nd	52	26
	S4	4	nd	nd	nd	nd	96	48	nd	36	12

Q = quartz, F_{Ca-Na} = calcium- and sodium-rich feldspars, F_K = potassium-rich feldspars, G = goethite, H = haematite, Ph = phyllosilicates, V = vermiculite, S = smectite, I = illite, K = Kaolinite, nd = not detected, tr = traces; -, not determined

Table 3.
Parameters of Langmuir isotherm at a reaction time of 1440 min.

Soil No.	Metal sol.	Q^a	CI^b	k	CI	SE^c	R^{2d}
		$\mu\text{mol kg}^{-1}$		reciprocal	$\mu\text{mol dm}^{-3}$		
S1	Pb	2.92×10^3	3.35×10^2	0.220	0.081	1.10×10^2	0.99
	Cd	2.37×10^3	2.25×10^2	0.002	0.001	0.97×10^2	0.98
	Pb(+Cd)	2.85×10^3	3.01×10^2	0.204	0.090	1.05×10^2	0.96
	Cd(+Pb)	8.03×10^3	1.86×10^2	0.006	0.003	0.63×10^2	0.96
S2	Pb	1.29×10^4	2.30×10^3	0.062	0.020	1.62×10^3	0.91
	Cd	7.67×10^3	1.57×10^3	0.003	0.001	6.39×10^2	0.95
	Pb(+Cd)	9.88×10^3	1.95×10^3	0.083	0.031	1.01×10^3	0.93
	Cd(+Pb)	3.36×10^3	8.18×10^2	0.006	0.003	2.80×10^2	0.95
S3	Pb	2.33×10^3	3.63×10^2	0.036	0.018	1.04×10^2	0.99
	Cd	1.27×10^3	6.99×10^2	0.002	0.000	0.71×10^2	0.96
	Pb(+Cd)	2.83×10^3	6.09×10^2	0.012	0.002	2.44×10^2	0.96
	Cd(+Pb)	1.86×10^3	4.58×10^2	0.001	0.000	0.66×10^2	0.96
S4	Pb	2.52×10^4	3.05×10^3	0.006	0.005	1.91×10^3	0.95
	Cd	1.69×10^4	1.08×10^3	0.001	0.000	2.05×10^2	1.00
	Pb(+Cd)	1.50×10^4	2.06×10^3	0.013	0.015	1.58×10^3	0.90
	Cd(+Pb)	7.29×10^3	1.14×10^3	0.003	0.002	3.67×10^2	0.97

^a Q is the maximum sorption capacity, k the bonding energy coefficient.

^b CI, 95 % confidence intervals of the estimated parameters

^c Standard error of estimate

^d All coefficients of determination were significant at a $P \leq 0.01$

Table 4.
Sorption - pH functional relationships ($\ln S = mpH + b$)

Metal sol.	Soil No	pH range ^a	<i>m</i>	CI ^b	<i>b</i>	CI	SE ^c	<i>R</i> ^{2 d}
Pb	S1	4.5-3.7	-2.83	0.51	11.63	2.09	0.124	0.82
	S2	4.1-3.7	-7.92	2.39	32.57	9.50	0.302	0.97
	S3	4.6-3.7	-3.40	1.13	14.56	4.62	0.296	0.83
	S4	4.1-3.6	-4.49	1.11	19.62	4.26	0.150	0.98
Cd	S1	4.1-3.7	-6.17	1.64	25.59	6.38	0.196	0.99
	S2	4.2-4.1	-27.38	9.31	115.64	38.41	0.236	0.96
	S3	4.0-3.6	-6.23	2.06	25.67	7.84	0.220	0.98
	S4	4.0-3.9	-11.62	2.91	49.24	11.60	0.072	0.99
Pb(+Cd)	S1	4.0-3.5	-6.62	1.61	25.99	6.09	0.213	0.90
	S2	4.3-3.6	-3.89	1.07	16.27	4.24	0.283	0.98
	S3	3.9-3.5	-7.56	2.13	30.03	7.99	0.287	0.98
	S4	3.9-3.5	-4.26	0.65	17.96	2.41	0.090	0.99
Cd(+Pb)	S1	4.0-3.5	-5.98	0.88	24.63	3.32	0.116	0.98
	S2	4.3-3.6	-3.36	1.20	15.56	4.75	0.266	0.96
	S3	3.9-3.5	-4.93	0.54	20.85	2.03	0.073	0.99
	S4	3.9-3.5	-3.10	0.74	14.32	2.76	0.102	0.97

^a pH ranges of the equilibrated solutions from the lowest and highest metal initial concentrations.

^b CI, 95 % confidence intervals of the estimated parameters

^c Standard error of estimate

^d All coefficients of determination were significant at a $P \leq 0.01$

Table 5.
Apparent sorption rate coefficients for different initial concentrations ($S = k t^{\nu}$)

	C_0^a	S1				S2				S3				S4			
		ν	CI ^b	SE ^c	R^2 ^d	ν	CI	SE	R^2	ν	CI	SE	R^2	ν	CI	SE	R^2
Pb	1	0.001	0.000	0.10	0.97	0.002	0.000	8.20	0.93	0.004	0.001	0.59	0.95	0.0005	0.000	2.10	0.94
	2	0.001	0.000	0.30	0.98	0.005	0.001	18.96	0.97	0.006	0.001	1.18	0.98	0.0007	0.000	3.63	0.96
	3	0.005	0.001	2.58	0.98	0.013	0.003	91.18	0.94	0.020	0.004	14.30	0.95	0.0023	0.000	8.06	0.98
	4	0.019	0.002	13.81	0.99	0.033	0.005	184.33	0.98	0.051	0.005	28.58	0.99	0.0028	0.000	20.63	0.98
Cd	1	0.023	0.003	2.43	0.97	0.016	0.003	44.20	0.95	0.034	0.005	2.60	0.97	0.0013	0.000	1.68	0.99
	2	0.036	0.003	2.82	0.99	0.022	0.003	48.36	0.98	0.035	0.005	4.42	0.97	0.0014	0.000	3.18	0.98
	3	0.042	0.004	10.12	0.98	0.030	0.005	91.62	0.97	0.035	0.006	8.33	0.97	0.0022	0.000	9.14	0.97
	4	0.055	0.008	30.60	0.98	0.057	0.006	126.82	0.99	0.031	0.006	14.81	0.96	0.0074	0.001	40.22	0.98
Pb(+Cd)	1	0.001	0.000	0.16	0.89	0.001	0.000	2.36	0.97	0.003	0.000	0.18	0.98	0.0003	0.000	1.04	0.91
	2	0.003	0.001	0.70	0.95	0.004	0.001	11.90	0.96	0.006	0.001	0.68	0.98	0.0007	0.000	2.46	0.94
	3	0.004	0.000	1.21	0.98	0.006	0.001	28.55	0.94	0.020	0.002	5.71	0.98	0.0011	0.000	4.71	0.96
	4	0.009	0.001	5.27	0.98	0.029	0.006	177.35	0.95	0.039	0.006	25.38	0.97	0.0030	0.001	23.93	0.93
Cd(+Pb)	1	0.023	0.003	0.76	0.98	0.020	0.003	17.55	0.97	0.040	0.005	0.71	0.98	0.0011	0.000	2.92	0.84
	2	0.031	0.001	0.53	1.00	0.038	0.005	35.38	0.98	0.052	0.007	2.28	0.98	0.0019	0.001	5.86	0.87
	3	0.059	0.009	7.85	0.97	0.050	0.010	74.35	0.96	0.054	0.009	5.43	0.97	0.0056	0.002	22.30	0.88
	4	0.063	0.008	9.19	0.98	0.054	0.010	85.24	0.96	0.054	0.007	6.30	0.98	0.0071	0.002	37.65	0.88

^a Increasing initial concentration solutions used for kinetic experiments as shown in table 1.

^b CI, 95 % confidence intervals of the estimated parameters

^c Standard error of estimate

^d All coefficients of determination were significant a $P \leq 0.01$

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Table 6
Q parameter - equilibrium time functional relationships ($Q = a t^b$)

Soil No.	Metal Sol.	a	CI ^a	b	CI	SE ^b	R ^{2 c}
S1	Pb	2517	31.16	0.0180	0.0026	28.36	0.98
	Cd	1506	104.36	0.0578	0.0135	94.82	0.94
	Pb(+Cd)	2380	146.14	0.0324	0.0073	126.36	0.94
	Cd(+Pb)	626	14.66	0.0364	0.0047	12.77	0.98
S2	Pb	10305	80.88	0.0304	0.0016	69.68	1.00
	Cd	5567	169.77	0.0455	0.0060	150.53	0.98
	Pb(+Cd)	7832	189.63	0.0335	0.0049	164.31	0.97
	Cd(+Pb)	2183	124.44	0.0634	0.0110	114.38	0.97
S3	Pb	1600	68.73	0.0491	0.0085	61.37	0.96
	Cd	1144	6.54	0.0241	0.0012	5.47	0.99
	Pb(+Cd)	2114	75.11	0.0393	0.0071	65.80	0.96
	Cd(+Pb)	796	67.52	0.0220	0.0154	69.00	0.98
S4	Pb	23916	203.35	0.0084	0.0018	168.37	0.95
	Cd	15387	127.38	0.0120	0.0017	106.14	0.97
	Pb(+Cd)	14413	191.61	0.0069	0.0028	158.23	0.83
	Cd(+Pb)	6909	73.62	0.0084	0.0022	60.96	0.92

5 ^a CI, 95 % confidence intervals of the estimated parameters
6 ^b Standard error of estimate
7 ^c All coefficients of determination were significant at a $P \leq 0.01$

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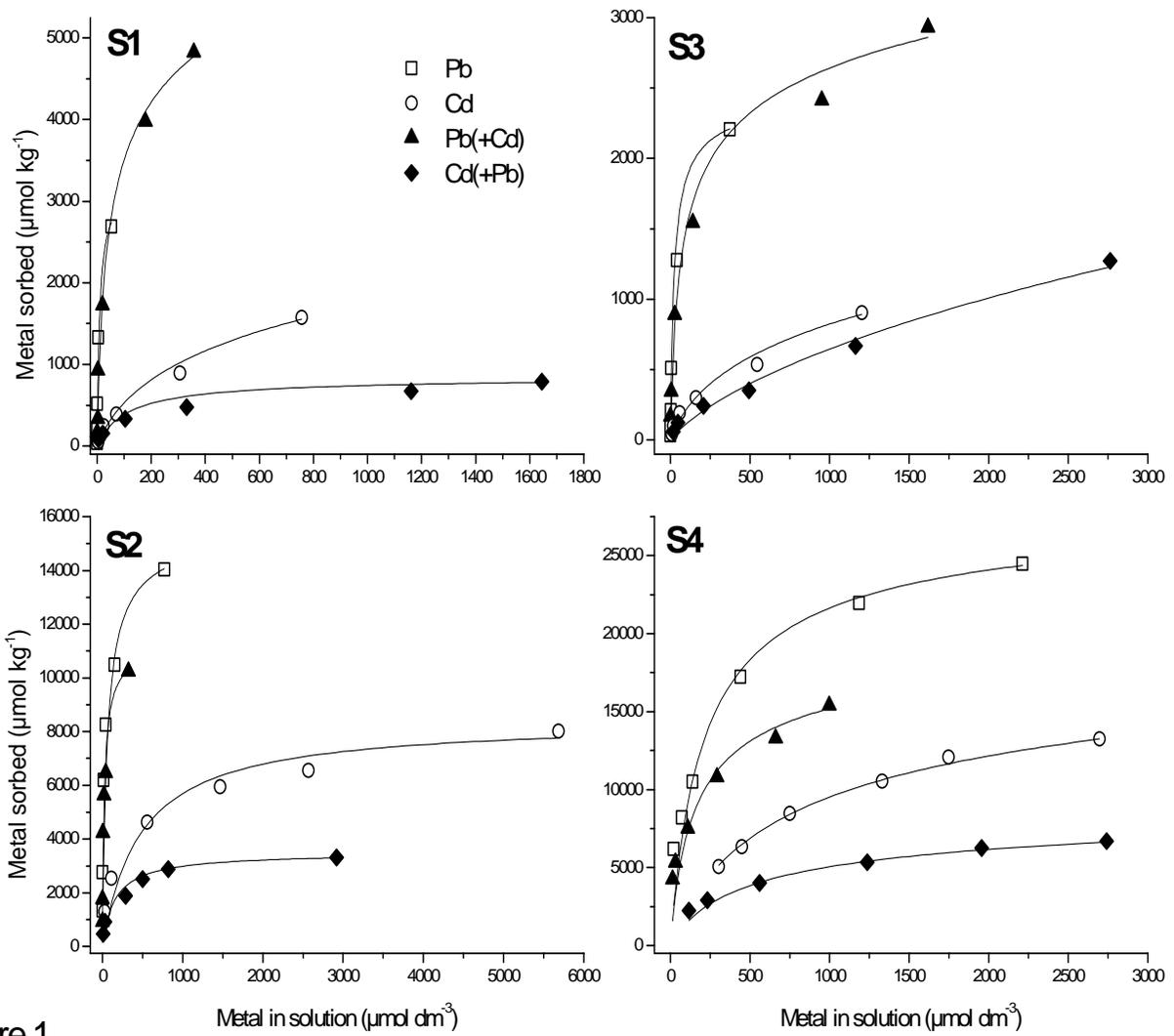


Figure 1

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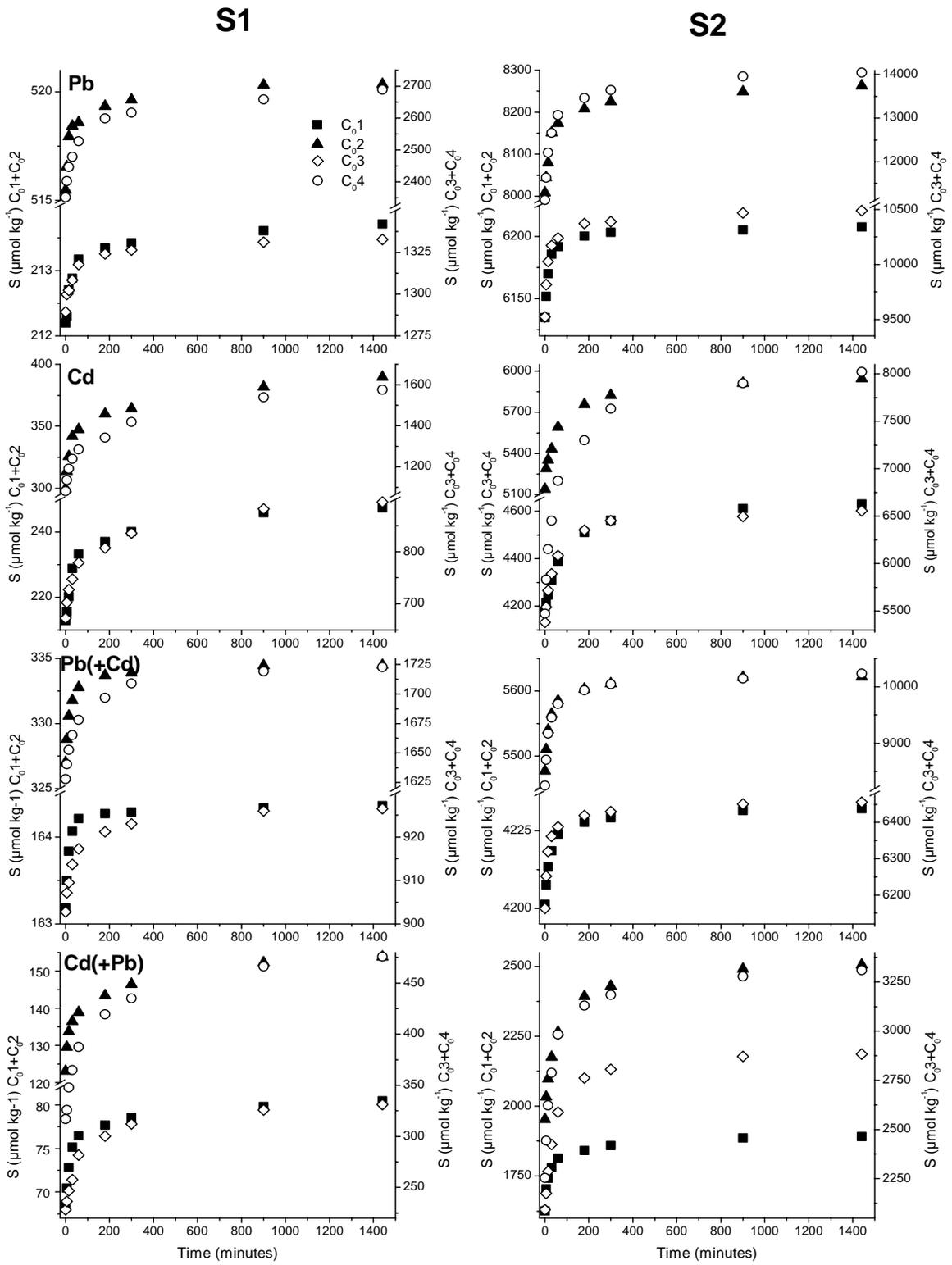


Figure 2